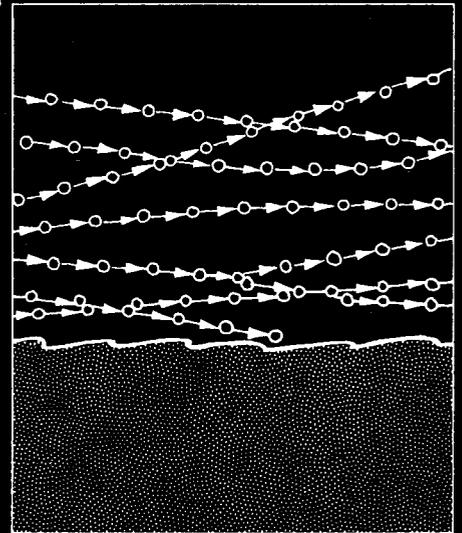


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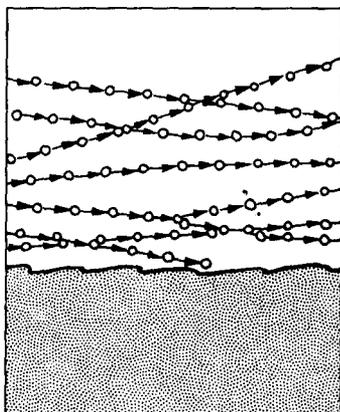
**A Water Resources
Technical Publication**

RESEARCH REPORT NO. 18

United States Department of the

INTERIOR

Bureau of Reclamation



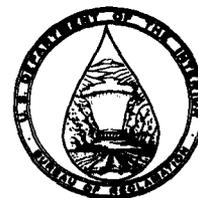
A study of the effects of waves on evaporation from free water surfaces

prepared by Calvin C. Easterbrook,
Cornell Aeronautical Laboratory, Inc.,
for the Bureau of Reclamation



UNITED STATES DEPARTMENT OF THE INTERIOR

BUREAU OF RECLAMATION



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PREFACE

For several years, the Bureau of Reclamation has conducted research in the use of monomolecular films (monolayers) of fatty alcohols and similar materials placed on water surfaces to retard evaporation. The incentive for this research is to reduce the great evaporation losses from large lakes and reservoirs in the 17 western United States, estimated at more than 14 million acre-feet annually, and at the same time to improve the quality of the water in the lakes and reservoirs.

The Bureau's research program in evaporation reduction is being carried out at selected field sites and in the Bureau's Engineering and Research Center at Denver, as well as through contracts with educational institutions, other government agencies, and commercial research organizations. Among these activities is the investigation to determine the quantitative relationship between wave and air flow characteristics and the evaporation rate, and to increase understanding of the microphysical processes that control the rate of mass transport across a water surface. Such an investigation, described in this Research Report, is important because, as the author states, "...waves

on a water surface have a significant effect on the evaporation rate from that surface".

The work described herein was carried out by Calvin C. Easterbrook of the Cornell Aeronautical Laboratory, Inc., of Cornell University at Buffalo, N. Y., under research contract No. 14-06-D-5764 with the Bureau of Reclamation. The original report was given limited distribution under its designation, CAL Report No. RM-2151-P-1, dated April 15, 1968. As Mr. Easterbrook's research work is of continuing interest to researchers studying the use of monolayers, it is reprinted here as a Bureau of Reclamation Water Resources Technical Publication.

Included in this publication is an informative abstract with a list of descriptors, or keywords, and "identifiers". The abstract was prepared as part of the Bureau of Reclamation's program of indexing and retrieving the literature of water resources development. The descriptors were selected from the *Thesaurus of Descriptors*, which is the Bureau's standard for listing of keywords.

Other recent issues in the Water Resources Technical Publications group are listed on the inside back cover of this report.

ACKNOWLEDGMENTS

I wish to express sincere appreciation for the assistance and cooperation given by Dr. Bradford Bean, Mr. Ray McGavin, and several other members of the Radio Meteorology Section, Tropospheric Telecommunications Laboratory, of the Environmental Science Services Administration. Without the meteorological data supplied by the ESSA group, the CAL study of wave-evaporation interaction at Lake Hefner would not likely have been possible. I also wish to thank the many people from the Bureau of Reclamation for their continued interest in and

support of this work and for their willing assistance during the Lake Hefner field program.

This report has been approved by George E. McVehil, Head, Dynamic Meteorology Section, and Roland J. Piliè, Assistant Head, Applied Physics Department of the Cornell Aeronautical Laboratory, Inc.

CALVIN C. EASTERBROOK.

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SUMMARY

Experiments performed in the laboratory and measurements made in the field show that waves on a water surface have a significant effect on the evaporation rate from that surface.

The laboratory measurements were made using a large wave tank-wind tunnel combination in which both wind speed and wave parameters could be controlled. Waves were generated at one end of the tank by a hydraulically driven paddle, and dissipated at the other end on a gradually sloped "beach." In this way, wave conditions approaching those developed on large bodies of water (fetches up to a mile or more) could be duplicated.

Results of the tank experiments indicated that for certain combinations of wind speed and well-developed waves, the evaporation rate was less than that measured under similar wind speed conditions with no waves present. Subsequent investigation of air flow over the waves showed that significant changes occurred as wave heights were increased. Regions of dead air were observed to the lee of the

wave crests, and vortices appeared in the wave troughs. The partially trapped air apparently forms an effective barrier to the vertical transport of water vapor.

As a follow-up to the laboratory work, wave and evaporation measurements were made at Lake Hefner in Oklahoma. Evaporation rates were computed from wind and moisture measurements by the eddy flux method, and correlated with wave data. The relationships so derived display characteristics very similar to those found in the laboratory.

The short-period mass transport data obtained during the Lake Hefner experiments were combined and adjusted to fit previously measured evaporation coefficient relationships derived from water and energy budget measurements. The resulting evaporation coefficient curve is believed to provide an improved measure of the "instantaneous" evaporation rate from Lake Hefner under given conditions of wind speed, water surface temperature, and ambient humidity.

INTRODUCTION

Many methods and equations have been derived over the past 100 years for estimating the amount of evaporation taking place from a free water surface. Some of these relationships are quite complex, requiring a number of precise meteorological measurements. Others, such as the simple mass transport equation investigated in this study, require only a few basic measurements, are easy to apply, and for these reasons have been widely used. Marciano and Harbeck¹ have shown that complex equations such as that of Sutton are little better in practice than the much simplified expressions which usually reduce to a form of the empirically derived mass transport equation,

$$E = NV_a(e_s - e_a) \quad (1)$$

where E = evaporation rate

V_a = mean wind speed at level a

e_s = saturation vapor pressure at temperature of the water surface

e_a = mean vapor pressure at level a

N = constant.

While equation (1) is a simple expression which apparently gives accurate estimates of evaporation amounts when applied over long periods of time, it can be defended on physical grounds as well. Equation (1) can be considered to be a simplified one-dimensional diffusion equation of the form

$$E = K \frac{\partial q}{\partial z} \quad (2)$$

where E = flux of q in the z direction

$\frac{\partial q}{\partial z}$ = gradient of q in direction z

K = diffusion coefficient.

In the free atmosphere, it is known that eddy diffusion by turbulence is much greater than molecular diffusion. Thus, the coefficient in the diffusion equation becomes the eddy diffusion coefficient K_w

for water vapor transport. Since the gradient of q must be continuous near the water surface, we can write

$$\frac{\partial q}{\partial z} = \frac{\Delta q}{\Delta z} \quad (2a)$$

which must be correct at some particular (but in general unknown) height in the region Δz . Equation (2) can then be written

$$E = K_w \frac{\Delta q}{\Delta z} \quad (3)$$

where K_w is the value of the diffusion coefficient appropriate to the height where equation (2a) holds.

If we consider the layer of air between the water surface and some height z , we can then write

$$E = D(q_s - q_z) \quad (4)$$

where

$$D = \frac{K_w}{z}$$

D , then, is a measure of the eddy diffusivity for the layer $(0-z)$ and, as such, is dependent on the height z and the characteristics of the boundary layer turbulence. The primary factor which governs turbulence intensity near the surface is wind speed. If we postulate a linear dependence of D on wind speed V_z , we can express the eddy diffusion equation as

$$E = NV_z(q_s - q_z) \quad (5)$$

which is exactly the empirical form of the mass transport equation.

The more complex evaporation equations attempt to express D analytically as a function of the many factors which influence turbulence. However, to the best of our knowledge, no attempts have been made to relate D to wave conditions, which may have an important effect on turbulence structure over the water. There are several ways in which waves might affect evaporation. First, we might think that increasing waves may constitute an increasingly rough surface and thus increase the turbulent transport of

¹ References in this report are listed on page 41.

water vapor. Also, the existence of bubbles, breaking on the wave crests, might constitute a significant decrease in the surface resistance to mass flux, thus increasing evaporation. Associated with this latter mechanism is another which might be even more important; that is, the creation of a vapor source above $z=0$ due to spray blown off breaking waves.

It is also possible to envision an effect opposite to the first suggestion above: instead of waves making the surface aerodynamically more rough, just the opposite may be true. A suggestion along these lines has been proposed by Stewart. Stewart presents arguments to show that traveling and developing waves on a water surface may give rise to an organized air flow pattern near the surface. This organized flow, while being able to transport momentum, would act as a barrier to heat and mass transport. A further argument related to the above is suggested by a postulate put forward by Phillips. Phillips found that waves develop most rapidly by a resonance mechanism which occurs when a component of the surface pressure distribution moves at the same speed as the free surface wave with the same wave number. Phillips' results suggest that certain combinations of wave conditions and wind speed might favor the development of organized flow as discussed by Stewart.

Because of the many unknowns in the evaporation process and its relation to waves on the evaporating surface, Cornell Aeronautical Laboratory undertook

a combined laboratory and field study of wave effects on evaporation for the Bureau of Reclamation. The study began in September 1965 and continued through January 1968 under two extensions to the original contract. The overall objectives of the work were as follows:

1. Determine the quantitative relationship between wave characteristics, air flow characteristics, and evaporation rate.
2. Investigate the microphysical processes that control the rate of mass transport across the water surface and the way in which these processes change with wave activity.
3. Develop methods for applying laboratory results to field situations.
4. Develop a general evaporation prediction technique and formulate simple expressions that make use of data from routinely available observations to obtain an estimate of the evaporation loss from untreated open water surfaces.

The work has progressed in two separate phases. Phase 1, completed in July 1966, consisted entirely of laboratory measurements and was directed mainly toward objectives 1 and 2. Phase 2 was a field study of evaporation processes at Lake Hefner, and was aimed at satisfying objectives 3 and 4. This document is the final report under the contract and describes the results of all work performed during the contract period.

THE WAVE TANK EVAPORATION EXPERIMENTS

The laboratory experiments described herein were carried out in a wave tank, originally constructed at CAL for use in the study of water waves by Doppler radar. The wave tank consists of a large open trough, 40 feet long, 4 feet wide, and 3 feet deep with an electrically driven paddle at one end of the tank. The paddle generates waves which travel down the trough and are dissipated on a gently sloping "beach".

For the evaporation experiments the wave tank was enclosed and a large fan was installed to circulate air over the water surface. Instruments were installed to make measurements of the moisture content, temperature, and speed of the air moving over the evaporating surface. Details of the evaporation experiments are given below.

Evaporation Tank Design and Instrumentation

A diagram of the combination wave and evaporation tank is shown in figure 1. A return duct constructed on top of the existing wave tank carries the

output air back to the input forming a closed recirculating system. A high capacity, variable speed fan installed in the return duct provides the controlled wind speed along the evaporation section of the tank. The ports at the beach end of the structure permit the system to be charged with dry air at the beginning of each experiment. Closing port *C* and opening ports *A* and *B* allows dry air to be drawn in through *A* while the tank is being purged of moist air through *B*. Wave generation is accomplished in the tank with a motor driven paddle, as shown in figure 1. The amplitude and frequency of the paddle excursion can be controlled in the paddle drive system, thus setting the desired wave parameters. Details of the paddle drive may be seen in figure 3. The evaporation tank in its final stages of construction is shown in figure 2.

During initial trial experiments with the completed tank it became apparent that the "beach" was introducing some undesirable effects. Waves which were stable down the main section of the tank began breaking on approaching the beach. Increasing the

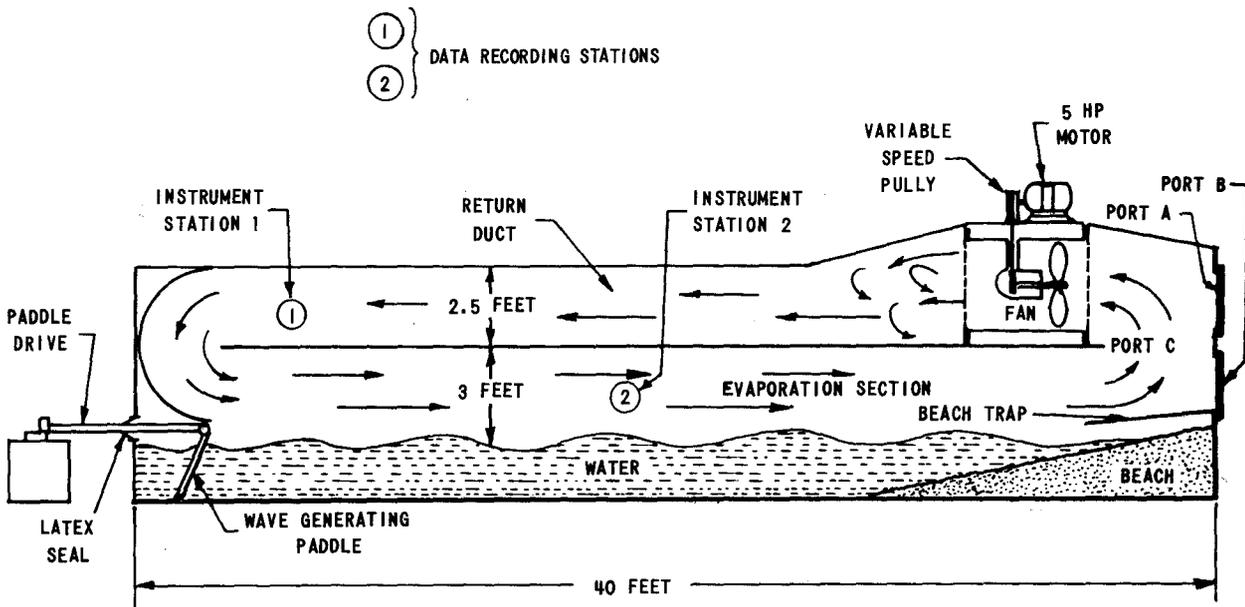


Figure 1.—Schematic diagram of evaporation tank.

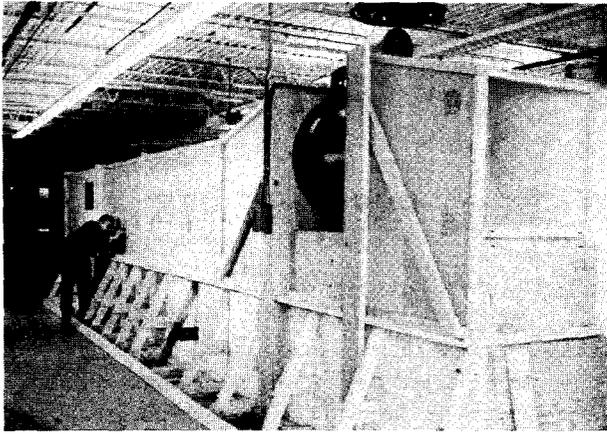


Figure 2.—Evaporation tank in final stages of construction.

wave height increased the beach action and hence increased the effective evaporation area. It was decided then to reduce the variable evaporation effects over the beach by creating a dead-air space over that region. Figure 6 shows the “beach trap” modification to the evaporation tank.

The trap was constructed using a polyethylene sheet stretched across the tank over the beach and sealed along the sides and the end. Some air exchange undoubtedly occurs at the mouth of the trap; how-

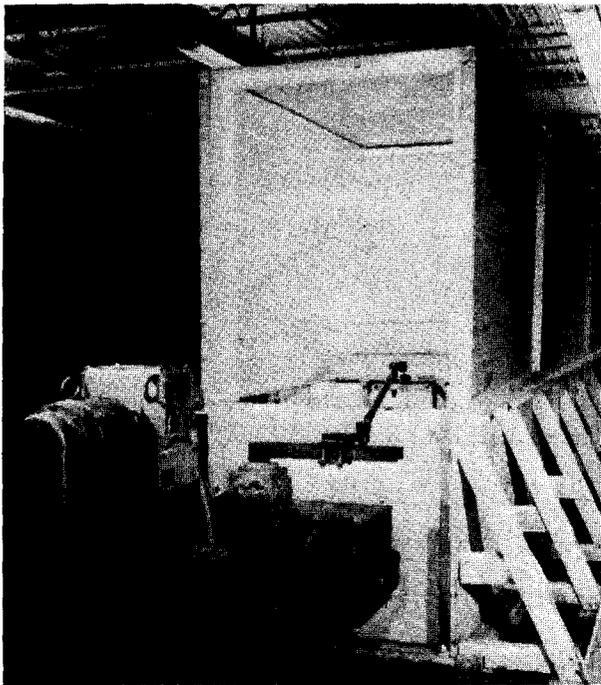


Figure 3.—Wave generating assembly.

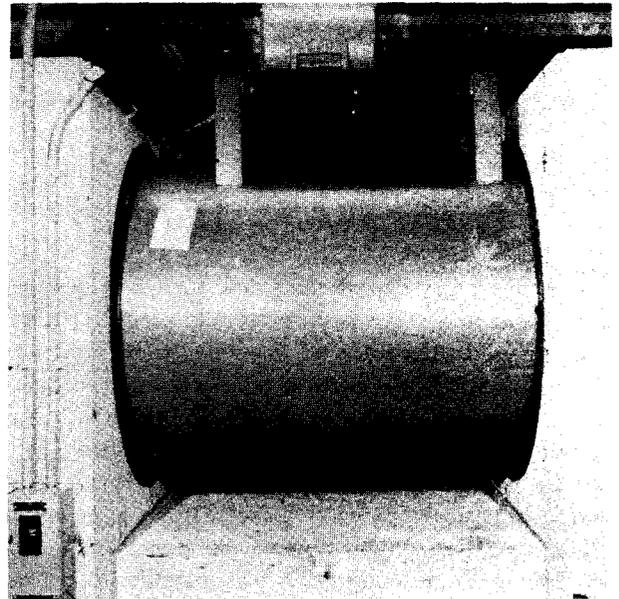


Figure 4.—High capacity fan installation.

ever, condensation taking place on the underside of the plastic sheet, even during the dry air charging cycle, is an indication of the high humidity that builds up there and demonstrates the effectiveness of the modification.

The temperature and humidity measuring instruments were mounted in the return duct at station 1

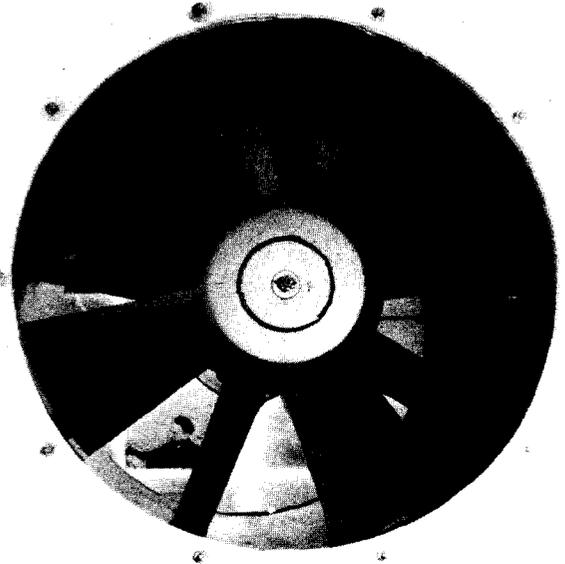


Figure 5.—The return duct observed through fan blades.

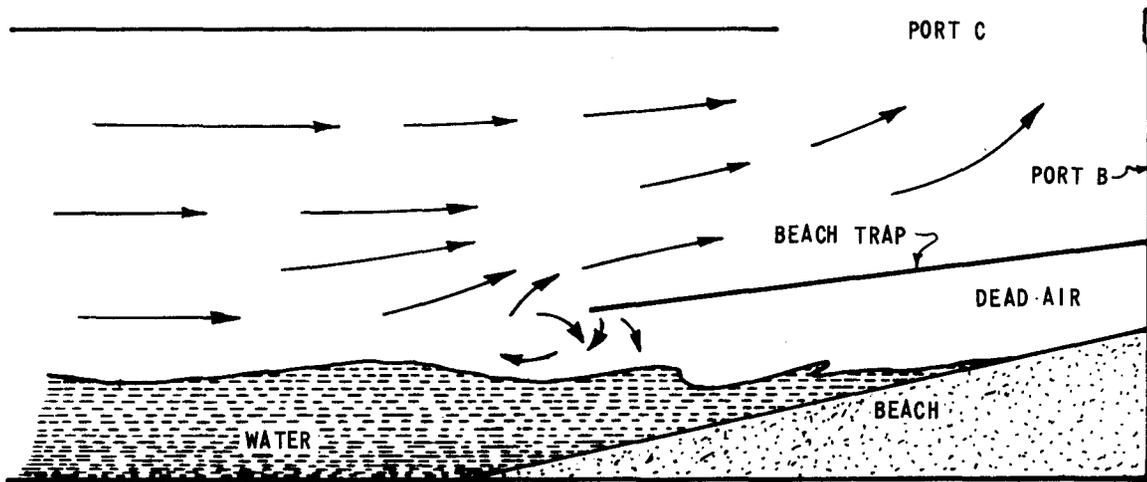


Figure 6.—Beach trap modification.

(figure 1). Temperature sensing was done with very small thermistors having a thermal time constant of approximately one-half second. The wet-bulb sensing thermistor was imbedded in a damp wick having a gravity feed. Approximately 2 inches of wick were exposed to the ventilating air. This permitted pre-cooling of the water before its arrival at the thermistor position and tended to decrease the response time for the wet-bulb measurement. The wet-bulb assembly was sheltered from the main return duct flow and was ventilated with a small auxiliary fan. The ventilation rate remained fixed at about 1,500 feet per minute. The time constant of the wet-bulb assembly was carefully measured and found to remain very nearly constant at 10 seconds for the full range of wind speeds in the return duct. The time constant of the air temperature measuring thermistor was of the order of 2 seconds.

Two special requirements of the temperature sensing instrumentation led to the use of a novel bridge circuit. First, the small thermistors required to get fast response have a very low thermal capacity and are easily heated by small electric currents, thus leading to false readings. Secondly, the evaporation experiments require a rapid sequence of simultaneous measurements of dry-bulb and wet-bulb temperatures. The bridge that was used was designed to dissipate only a few microwatts in the thermistor and to permit rapid data taking. A schematic diagram and a description of the bridge are given in appendix A.

The wind measurements were made with a hot-wire anemometer installed at station 2 in the main

duct, about midway between the water surface and the top of the duct. Also mounted in the water at station 2 was a mercury thermometer for measuring the bulk water temperature. This thermometer was completely emersed about 2 inches off the bottom of the tank and was read through the transparent tank wall at this position.

Development of Evaporation Tank Theory

In the evaporation experiments, room air is initially drawn into the tank, circulated over the water surface, and exhausted outside the building. When an equilibrium has been established for the desired wind and wave parameters (as indicated by constant values of wind speed and humidity) the tank is sealed and the air circulates over the water and through the return duct in a closed system. The rate of change of humidity with time is then recorded. The evaporation rate can be computed from the observed changes in humidity; however, it is possible to measure the evaporation coefficient directly with the system, and this approach was utilized to provide a single measure of evaporation rate for each desired combination of wind speed and wave state.

Let E represent the effective evaporation in grams per centimeters squared per second averaged over the whole tank water surface of area A at any particular time t . The total evaporation from the tank at this time is then EA grams per second, which is added to the total air volume v in the tank. Thus, the rate of change of specific humidity, as measured at station 1 when the system is closed (assuming com-

plete mixing in the return duct), is

$$\frac{dq}{dt} = \frac{AE}{v\rho} \text{ grams per gram per second} \quad (6)$$

or

$$E = \frac{v\rho}{A} \frac{dq}{dt} \quad (7)$$

where ρ = density of air

q = specific humidity.

Equation (7) then permits a computation of evaporation rate from the slope of the measured $q(t)$ curve if the air volume in the tank and the effective water surface area are known.

Let us next consider the simple bulk aerodynamic evaporation equation

$$E = \rho N (q_s - q_a) V_a \quad (8)$$

where E = evaporation rate, grams per centimeter squared per second

ρ = air density, grams per centimeter cubed

q_s = specific humidity of saturated air at the temperature of the water surface, grams per gram

q_a = specific humidity at some fixed level a

V_a = wind velocity at level a

N = evaporation coefficient.

The evaporation coefficient N is normally treated as a constant but it is known to be dependent upon atmospheric stability and, as shown in our preliminary experiments, is dependent upon wave structure. Furthermore, equation (8) breaks down when the wind velocity approaches zero, since evaporation does not stop entirely under zero wind velocity provided that $q_a < q_s$. We elect, therefore, to incorporate the wind velocity dependence into the factor D . D now has units of velocity and is a function of V as well as other parameters describing the wave characteristics and turbulent transport. Equation (8) then may be written in the form

$$E = \rho D (q_s - q_a) \quad (9)$$

In the wave tank experiment, equation (9) can only be applied on an instantaneous basis since q_a and E change with time. Thus, for the tank experiment we can write equation (9) as

$$E(t) = \rho D [q_s - q(t)] \quad (10)$$

Here, $q(t)$ is the instantaneous value of q measured at station 1 in the tank. If we combine equations (7)

and (10) and differentiate with respect to t , we obtain the differential equation

$$\frac{v}{A} \frac{d^2q}{dt^2} + D \frac{dq}{dt} = 0 \quad (11)$$

This result is obtained under the assumptions that q_s is constant (water temperature remains constant) and D is a constant for a particular experiment. Equation (11) has the solution

$$q(t) = q_s \left[1 - \exp\left(-\frac{DA t}{v}\right) \right] \quad (12)$$

for the boundary conditions of our experiment, i.e.,

$$q = 0 \quad \text{for } t = 0$$

$$q = q_s \quad \text{for } t = \infty$$

The quantity $\frac{v}{DA}$ is the time constant of the evapora-

tion experiment under the particular wave and air velocity conditions set and we can write

$$D = \frac{v}{A\tau} \quad (13)$$

where v = effective volume of air in tank

A = effective evaporation area

D = evaporation coefficient (with units of velocity)

τ = time constant of experiment.

Thus, D can be computed directly from measurement of the experimental time constant.

There are two sources of error inherent in the actual tank apparatus that are not accounted for in the simple solution above. First, the humidity measuring instruments at station 1 do not respond instantaneously, but have a finite time constant as discussed in the preceding section on instrumentation. Also, the tank is not completely air tight. There is a small amount of air leakage around the fan bearings, the paddle drive seal, and the intake and exhaust ports. It is possible to take these effects into account in the final analysis of the tank experiment.

Considering the evaporation process and the resulting solution, equation (12), it is readily seen that the recirculating tank experiment is directly analogous to charging a capacitor through a resistance. Here the time constant v/DA is equivalent to RC in the electrical circuit, R is equivalent to $1/DA$, and C is represented by the tank volume v . With this in mind we can set up an electrical analog of the evapo-

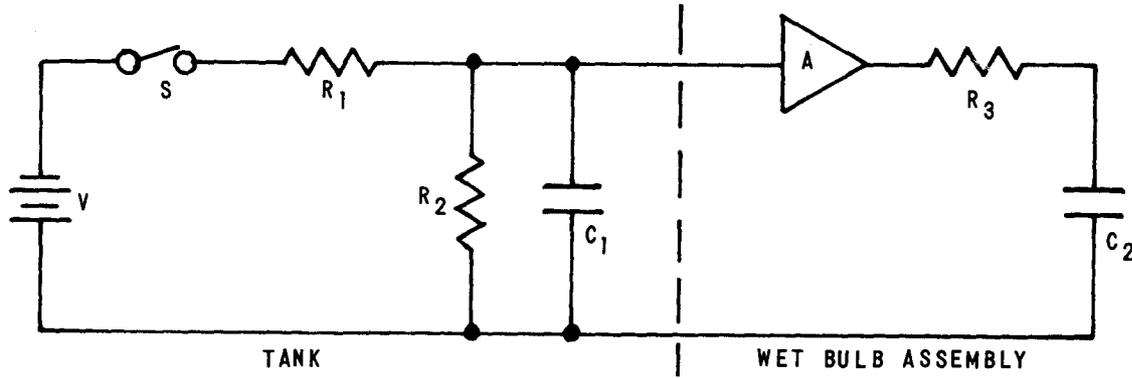


Figure 7.—Electrical analog of evaporation experiment.

ration tank complete with air leakage and measuring instrument time constant. This analog is shown in figure 7. The parallel resistance R_2 represents the tank leakage and R_3C_2 represents the time constant of the wet-bulb assembly. Amplifier A , which has unity gain and an infinite input impedance, simply isolates the measuring device from the driving circuit. The solution for this circuit is developed in appendix B. Applying the circuit solution to the evaporation experiment, it is shown in appendix B that the complete expression for the specific humidity as a function of time is given by

$$q = q_0 + (q_m - q_0) \left[1 - \frac{r\tau}{r\tau - \tau_m} \exp\left(-\frac{t}{r\tau}\right) + \frac{\tau_m}{r\tau - \tau_m} \exp\left(-\frac{t}{\tau_m}\right) \right] \quad (14)$$

where q_m = measured specific humidity at the end of an experiment

q_0 = measured q at $t = t_0$

q_s = saturation value of q at the water temperature

$r = q_m/q_s$

τ_m = time constant of the wet-bulb assembly

τ = desired time constant.

Considering the measured response time of the wet-bulb assembly and the approximate minimum experimental time constants it is found that the third term inside the square brackets in equation (14) is negligible for $t \lesssim 20$ seconds. The theoretical expression for the experiment then becomes

$$q = q_0 + (q_m - q_0) \left[1 - \frac{r\tau}{r\tau - \tau_m} \exp\left(-\frac{t}{r\tau}\right) \right] \quad (15)$$

Rearranging equation (15) and taking logarithms we

have

$$\ln \left[1 - \frac{q - q_0}{q_m - q_0} \right] = \ln \frac{r\tau}{r\tau - \tau_m} - \frac{t}{r\tau} \quad (16)$$

Equation (16) produces a straight line with slope $-\frac{1}{r\tau}$ when plotted on semi-log paper. The quantity

$r\tau$ is the actual measured time constant, which can easily be obtained from the semi-log plots. The time constant τ that would have existed without leakage into the tank may be computed using the measured value q_m and tabular values for q_s at the water temperature.

With the analysis of the evaporation experiment complete, it is then possible to set up the experimental procedure necessary to meet the specified boundary conditions. With reference to figure 1, the procedure is as follows:

1. A sliding "valve" is closed over port C.
2. Port A is opened to the room where the wave tank is located so that room air can be drawn into the tank.
3. Port B is opened and connected to a flexible duct, which discharges outside the building.
4. The fan is started and set for the required wind speed. Dry air is thus circulated through the system and exhausted outside the building.
5. Desired wave height and wave length are established by proper setting of the paddle drive controls.
6. Initial equilibrium is established at the humidity and temperature measuring station (station 1). This determines the initial value of humidity, q_0 .
7. Port C is then quickly opened and ports A and B closed. A stop watch is activated at this time, establishing t_0 .
8. The wet- and dry-bulb temperatures at station 1

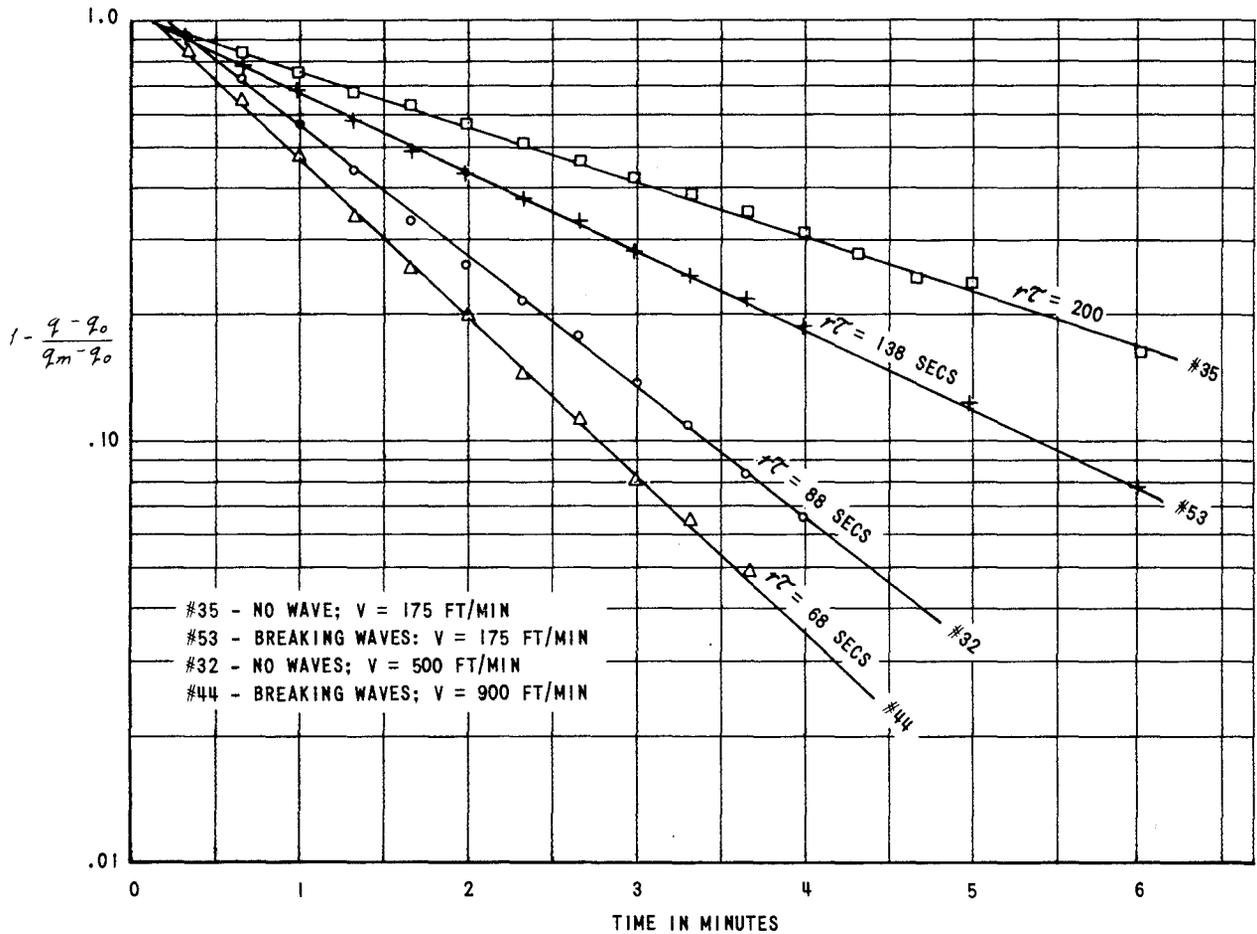


Figure 8.—Normalized specific humidity as a function of time.

are recorded every 20 seconds during the first few minutes of each experiment, and then less frequently as the evaporation rate falls off due to the increase in moisture content of the air. The maximum equilibrium value q_m is obtained when the wet-bulb temperature ceases to increase.

9. The average wind speed at station 2 during the experiment is recorded, along with the water temperature. The barometric pressure at the time of the experiment is also recorded.

10. The quantity $1 - \frac{q(t) - q_0}{q_m - q_0}$ is computed for each

data point and the values are plotted on semi-logarithmic graph paper. The slope of the resulting straight line then represents the time constant of the experiment. Note that the initial point q_0 is not used in determining the appropriate straight line through the data points. This is in keeping with the approximation used when neglecting the last term of the solution, equation (14).

Results of the Wave Tank Evaporation Experiments

A total of 108 experiments was completed in the wave tank, producing 76 useful samples of data. The first 32 data runs were performed during the "shake-down" period of the evaporation-wave tank while modifications were still being made; hence, these data were not used in the final analysis.

Semi-log plots of the data from several representative experiments are presented in figure 8. The fit of the data points to the theoretical straight line of equation (16) is very good. From each semi-log plot the experimental time constant $r\tau$ was measured and reduced to the idealized time constant τ by dividing by factor $r = q_m/q_s$. The evaporation coefficient D was computed using the relationship $D = v/A\tau$. The tank parameter v/A , computed from careful measurements of the inside dimensions of the tank, was found to be 5.8 feet. The measured time constants ranged from 38 to 218 seconds leading to a range of D

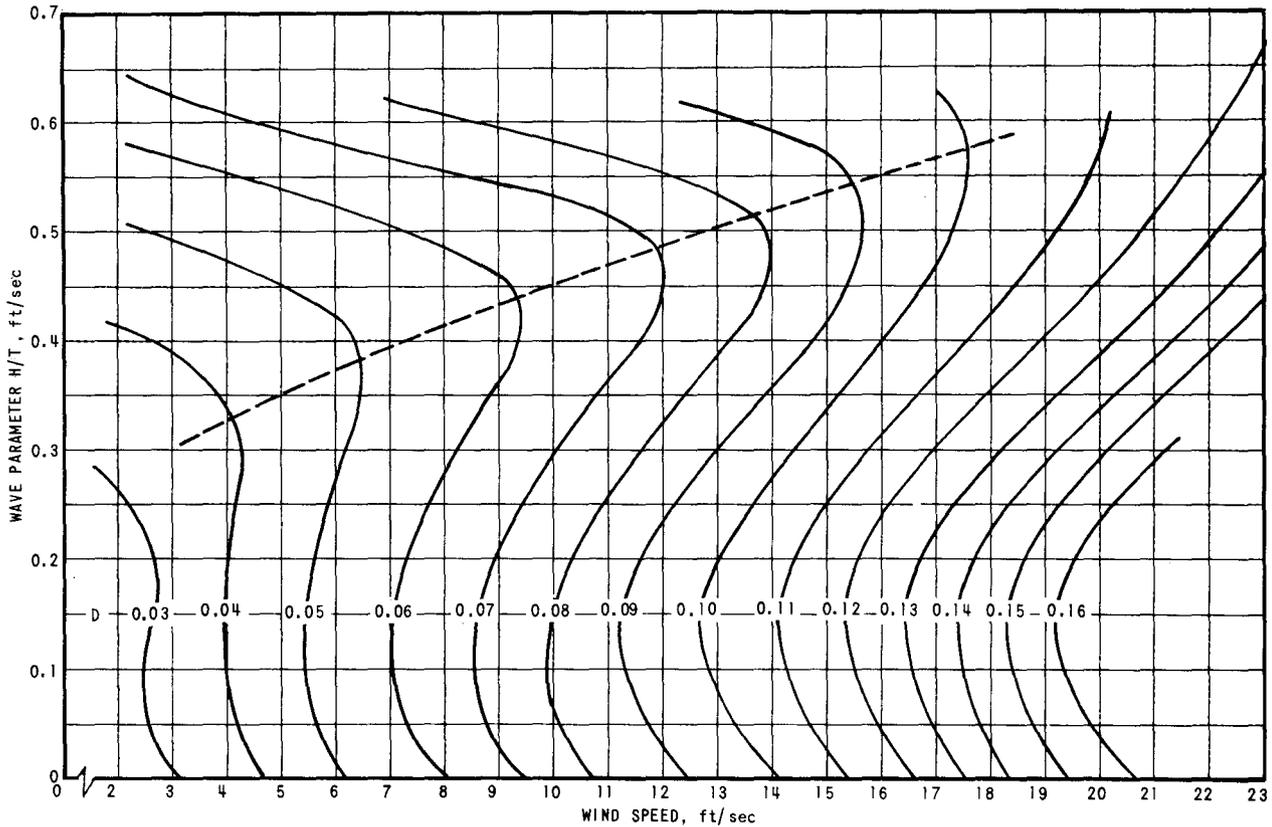


Figure 9.—Evaporation coefficient D in feet per second as a function of wind speed and wave parameter H/T .

factors from 0.027 to 0.152 foot per second. All basic data from the final 76 experiments were compiled and are included in appendix D of this report.

During the data analysis, two wave parameters were considered: the wave steepness parameter H/L (wave height to length ratio) and the wave height to period ratio H/T . Correlation of either of these two parameters with the evaporation data produced essentially the same results. We have chosen to present the final results in terms of the wave parameter H/T . This parameter seems more appropriate for two reasons: (1) wave period is easier to measure than wave length, and (2) the parameter H/T contains information on wave velocity as well as wave length, both of which may be important outside the laboratory.

Figure 9 shows the final results of the wave tank analysis. This diagram was constructed by plotting the measured values of the evaporation coefficient D at the appropriate location in the $H/T, V$ field and drawing contours of constant D . The resulting isopleths display several interesting characteristics. The first property one may note is that for a given value

of H/T the evaporation coefficient is a nearly linear function of wind speed. This result is not surprising in view of the success with which the empirical equation (8) has been applied in the past. Secondly, it is evident from figure 9 that D is not constant for a given wind speed, but is also a function of wave characteristics. It is here that a rather unexpected result appears. Over a certain range of wave characteristics, the evaporation rate actually decreases with increasing wave height to period ratio. (This result is discussed in some detail later.) Finally, there is a break-over zone (shown on figure 9 as a dotted line) above which the D factor increases rapidly with increasing H/T . This zone roughly coincides with the initial appearance of a visible curl on the wave crest, indicating wave breaking. The correspondence can be described only as approximate because the curl was never observed with H/T smaller than 0.5, but the evaporation data indicate that at low wind velocities the break-over zone does occur at values of H/T as small as 0.3. In previous work we have noted that other properties of waves in still air approach values characteristic of breaking waves before a

visible curl appears. In particular, Mee reported similar variations in the Doppler spectrum of sea clutter.

Having obtained evaporation data in the laboratory, it is of interest to compare the results with evaporation measurements made in the field. One sample of evaporation data with which we can compare our results is that obtained in the study of Lake Hefner for the Bureau of Reclamation by the Navy Electronics Laboratory (Marciano and Harbeck). The empirical equation of best fit to the data given in this report is

$$E = 6.25 \times 10^{-4} u_8 (e_0 - e_8) \quad (17)$$

where E = evaporation rate in centimeters per 3 hours

u_8 = 8-meter wind speed in knots

e_0 = saturation vapor-pressure at the water surface temperature in millibars

e_8 = vapor pressure at 8 meters.

The evaporation equation used in the CAL study is

$$E = \rho D (q_0 - q_a) \quad (18)$$

where E = evaporation rate in grams per centimeter squared per second

ρ = density of air in grams per centimeter cubed

q_0 = saturation specific humidity at water surface temperature

q_a = measured specific humidity at a point in the return duct of the wave tank.

If we use the relationship

$$q \approx \frac{0.622e}{p} \quad (19)$$

where p = atmospheric pressure in millibars

e = vapor pressure in millibars

and make the necessary transposition of units, we can reduce the tank evaporation equation to

$$E = 0.244 D (e_0 - e_a) \text{ centimeters per 3 hours} \quad (20)$$

For a wind speed of 10 knots and a wave parameter $H/T=0.4$, for example, the two equations agree within a factor of 4, the larger evaporation rate being given by the tank-developed equation. This result is not unreasonable if it is remembered that the N.E.L. equation uses parameters measured at 8 meters, whereas the CAL wave tank equation was derived from measurements made within one-half meter of the surface. The two equations cannot be reconciled precisely without detailed knowledge of the wind and humidity profiles both in the field and in the

wave tank. However, the factor of 4 can be accounted for using reasonable values of vertical wind shear and moisture gradient. A detailed comparison is not really justified since it is clear that air flow properties in the tank cannot fully duplicate those in the atmosphere. The experiments have been conducted to determine the variability of evaporation with wave state rather than to measure precise evaporation rates for application to field situations.

Investigation of Evaporation Controlling Mechanisms

The presentation of experimental results in figure 9 is appropriate for use in evaporation rate computations. The variation of evaporation coefficient D with wave characteristics is more obvious in a presentation of the same data showing D as a function of wave parameter H/T for constant values of wind speed as given in figure 10. An unexpected feature mentioned earlier is immediately evident in figure 10. A large region exists (shaded area in the figure) in which the evaporation coefficient actually decreases with increasing wave action and the inverse relationship becomes more pronounced at higher wind speeds. This type of behavior was not anticipated at the beginning of the study but appears to be a significant feature of the wave tank results. It is clear that an understanding of the mechanisms that control D must include an explanation of this phenomenon. The experiments described below were conducted in attempts to explain this behavior.

1. Surface temperature study

Upon completion of the evaporation measurements and having noted some of the peculiarities in the results, an attempt was made to determine the causes of the observed behavior. One evaporation retarding mechanism that must be considered is the possible reduction in surface temperature due to the evaporation itself. An infra-red thermometer was installed in the tank with the sensor viewing the water surface. The intent was to measure the water surface temperature under many different surface and evaporation conditions. However, no significant differences in temperature could be detected. Failure of the IR thermometer to detect surface temperature changes was not considered conclusive since it was thought that the instrument was being influenced by radiation from the walls and ceiling of the duct which was reflected into the sensor from the water surface.

While the simple IR thermometer installation failed to detect surface temperature changes, these changes were in fact detected by the duct thermistor

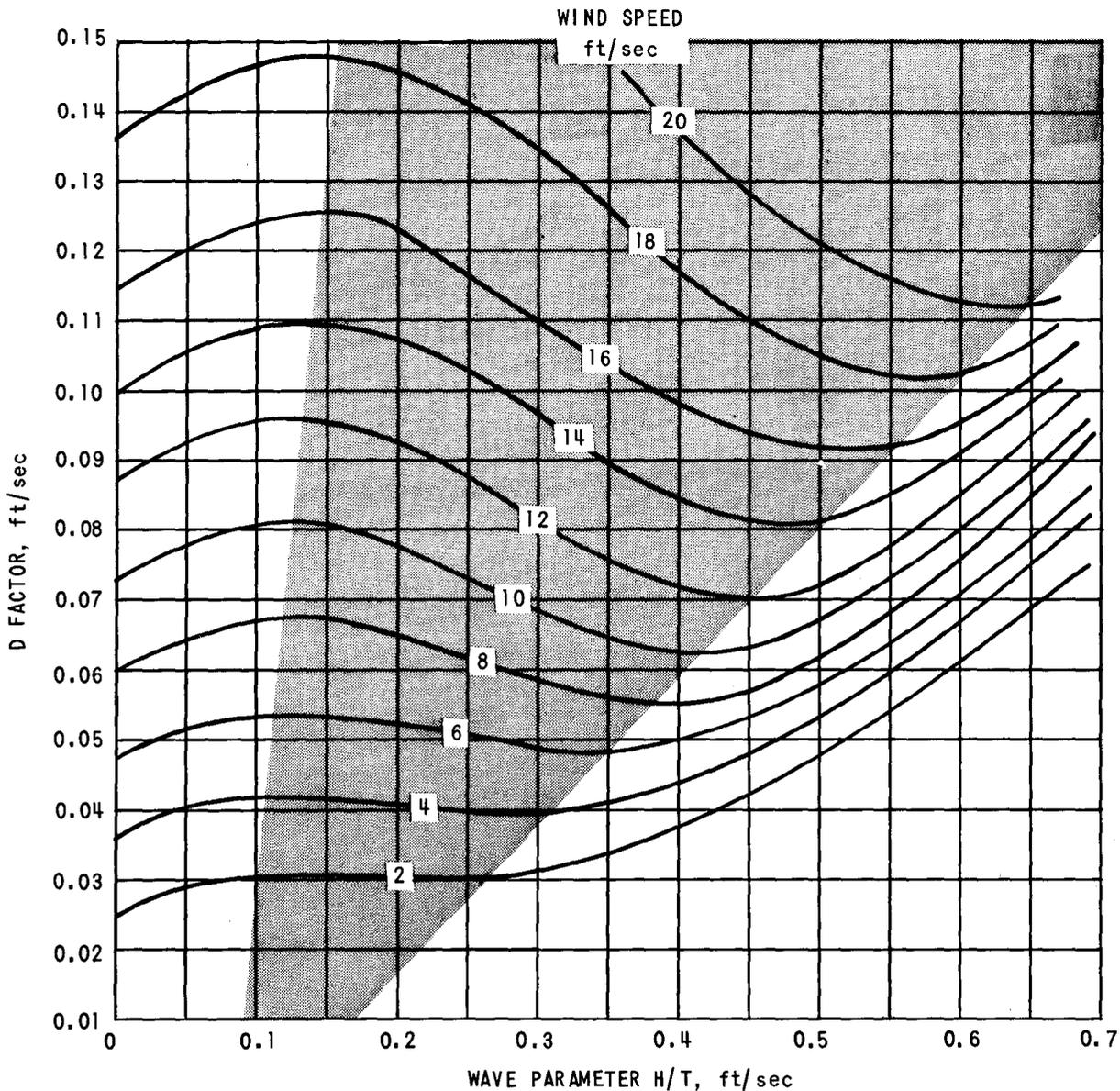


Figure 10.—Evaporation coefficient as a function of wave parameter for constant wind speeds.

used for the main temperature measurement. A temperature drop was noticed only during the high wind speed runs, where the initial evaporation rates were high enough to change the surface temperature by a detectable amount. Figure 11 shows an example of data from one of these experiments. The time constant represented by this data run was 41 seconds, giving a D factor of 0.142 foot per second. Note that during the first 20 seconds, when the evaporation rate is maximum, the air temperature fell a full degree and then gradually climbed back to the final equilibrium value of 24.7° C. The bulk water temperature during this experiment was 23.0° C. Thus,

the equilibrium temperature with no evaporation (end of experiment) was 1.7 degrees above the water temperature. The difference is due to the warming of the air by the walls of the return duct during the period over which the air leaves the evaporating surface and reaches the temperature measuring station. It is not unreasonable to assume that the heat transfer through the duct walls was higher at the beginning of the experiment where the inside to outside temperature difference was greatest. Thus, we can assume that at the 20-second point the measured air temperature inside the duct was at least 1.7 degrees above the integrated water surface

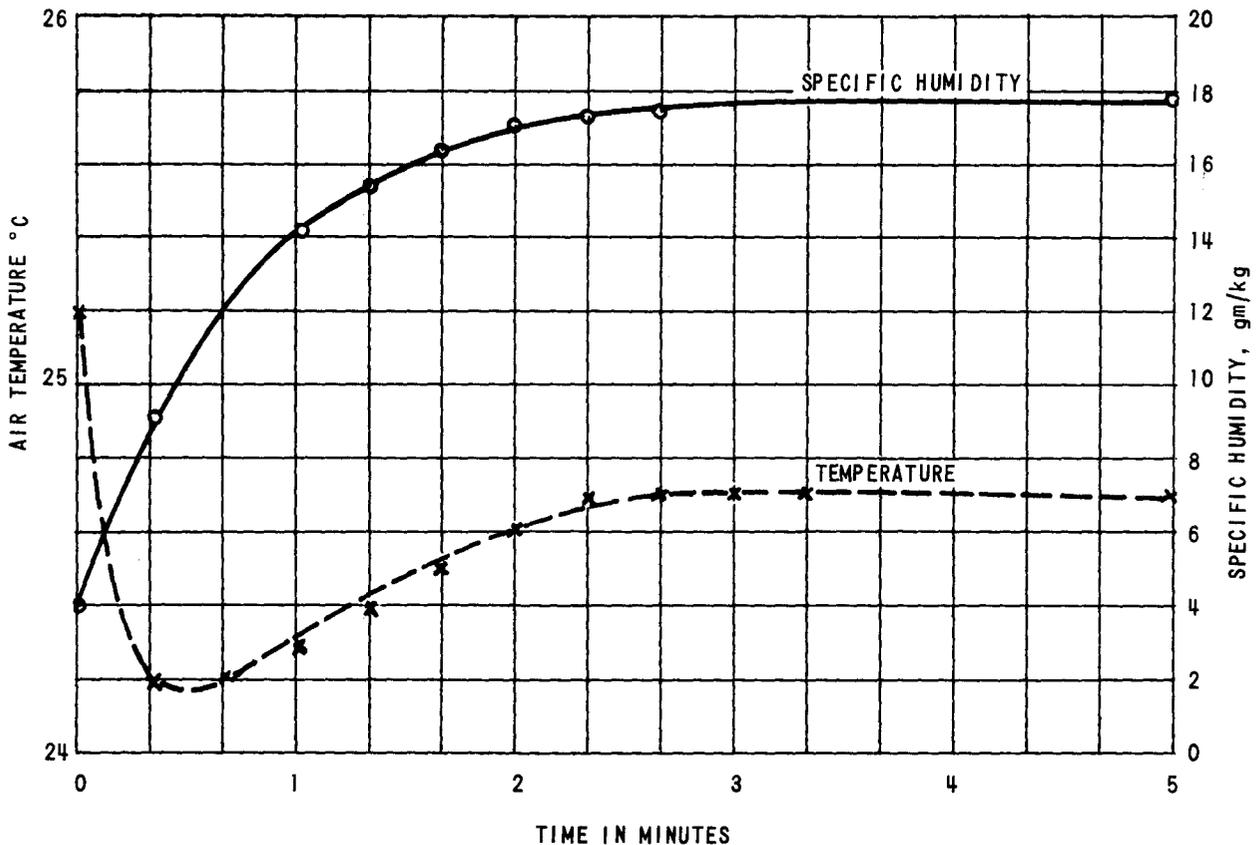


Figure 11.—Temperature and humidity as a function of time, Experiment No. 108.

temperature. This assumption is certainly on the conservative side, but it still places the water surface temperature at 22.5°C , or 0.5 degree below the bulk water temperature. Although the temperature drop is small and not likely to be significant in reducing evaporation rates, its presence is rather surprising in view of the violent water surface agitation by wind and waves taking place during the experiment.

2. Air flow tracing

Another possible reason for the decrease in evaporation rate with increasing wave steepness is the reduction of vertical mass transport due to modification of the air flow and turbulence characteristics. In order to investigate this postulate, apparatus was set up in the wave tank to inject tracers into the air stream above the water.

The first method used was to inject neutrally buoyant soap bubbles inflated with helium into the air flow above the water (Schooley). Movies were taken with a 16 mm camera running at approximately 50 frames per second. Displaying these pictures at 8 frames per second uncovered some interesting flow structure around well-developed waves.

Figure 12 is an example of several consecutive frames showing a bubble in a typical stationary position (relative to the wave) just in the lee of a wave crest. The same sequence shows other bubbles at higher altitudes moving with the wind speed, which was 2 to 3 times the wave phase velocity. Similar data were analyzed frame by frame, plotting bubble trajectories. Typical trajectories over small, wind-driven capillaries are shown in figure 13. Note the fairly uniform motion at all levels. Figure 14, in turn, shows bubble trajectories in the presence of well-developed waves. Note that the upper flow is relatively undisturbed but the low level circulation is completely changed. Circulation patterns coupled to the wave profile are immediately evident.

In another series of experiments, air motions were traced by releasing chemical vapor plumes near the waves and photographing the plumes with the camera. Air, saturated with titanium tetrachloride (TiCl_4), was pumped through the same bubble generators used in the bubble tests. TiCl_4 reacts with water vapor to form a fairly dense white cloud of hydrogen chloride. The fumes are toxic, of course, but the experiments were carried out in the sealed

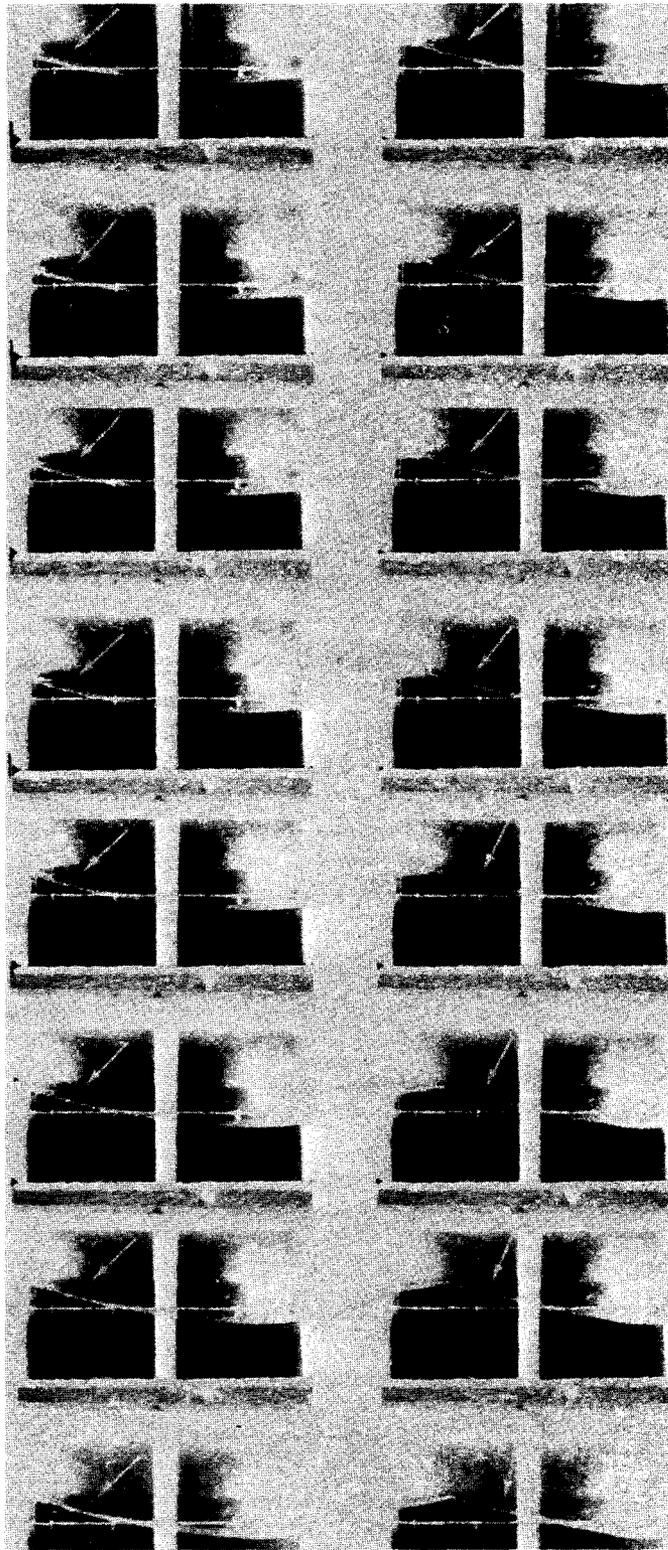


Figure 12.—Helium bubble stationary behind wave crest.

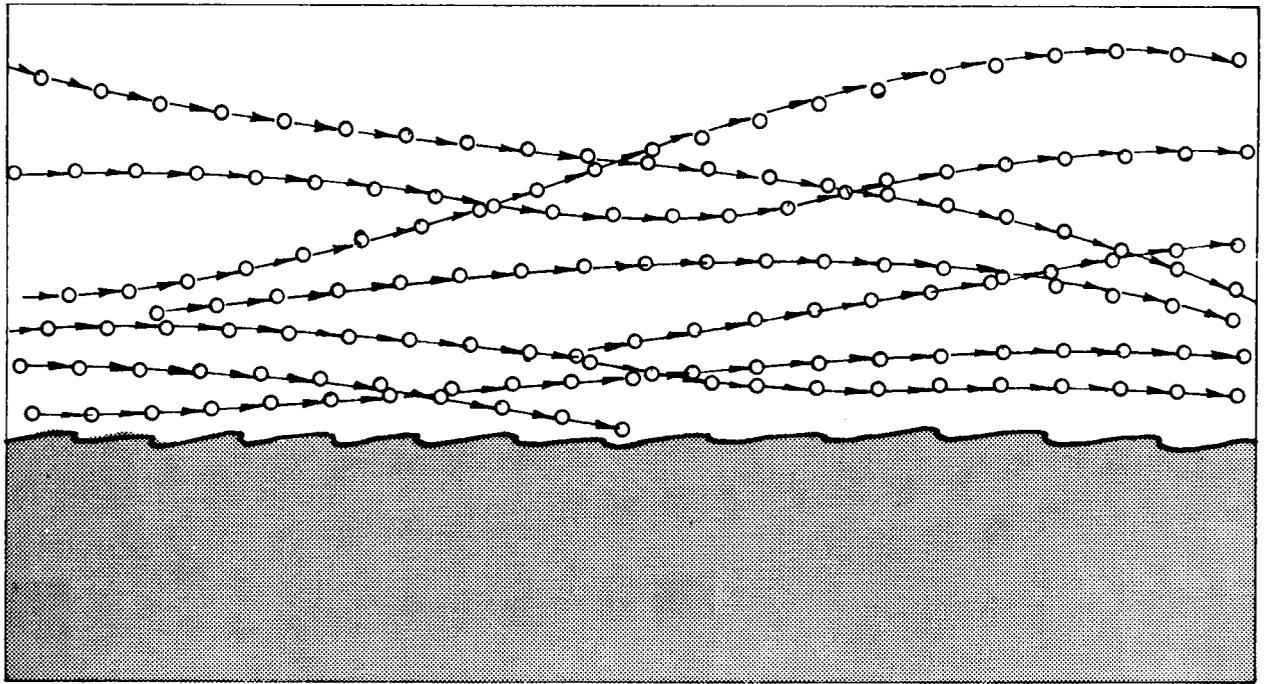


Figure 13.—Bubble trajectories over surface with only wind-driven capillary waves present.

evaporation tank and no particular difficulties were encountered. The tank environment was ideal for this type of tracer since the experiments could be run at almost 100 percent relative humidity, the condition necessary for maximum reaction of the $TiCl_4$.

Figures 15, 16, 17, and 18 show typical frames from this series of pictures. In figure 15a, the vapor plume traces an upward flow as the wave crest approaches. This flow continues to the front edge of the wave (figure 15c), then shifts to downward flow on

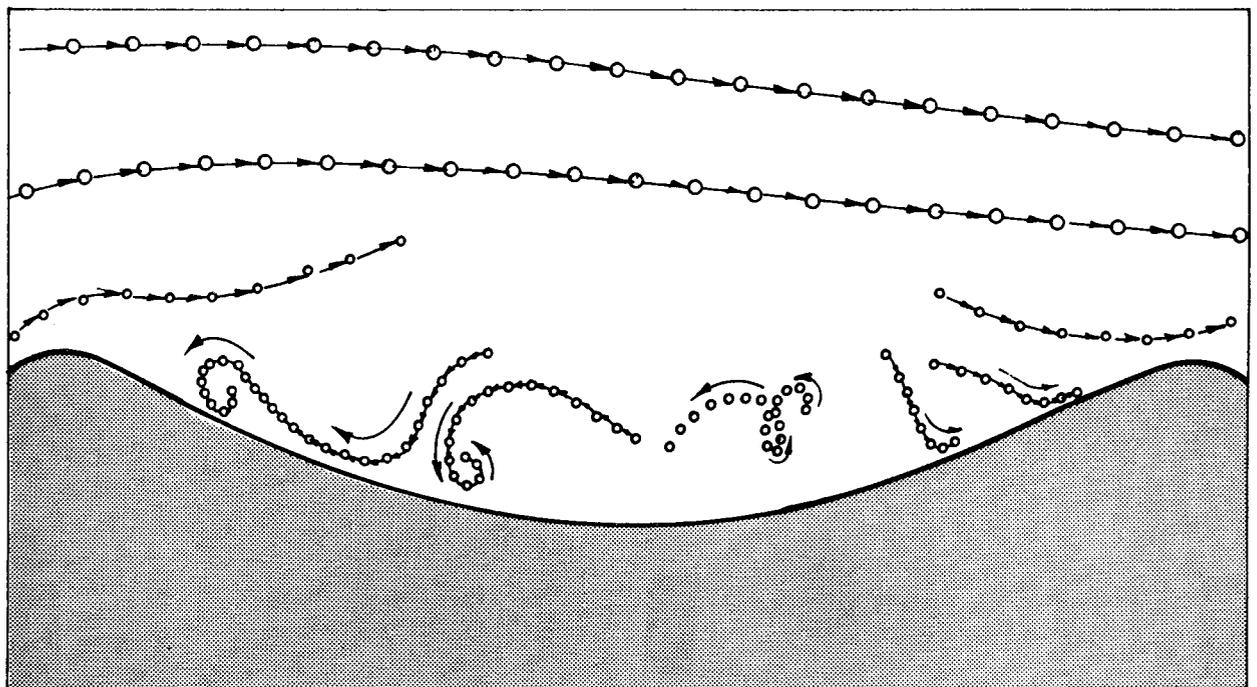


Figure 14.—Bubble trajectories over well-developed waves.

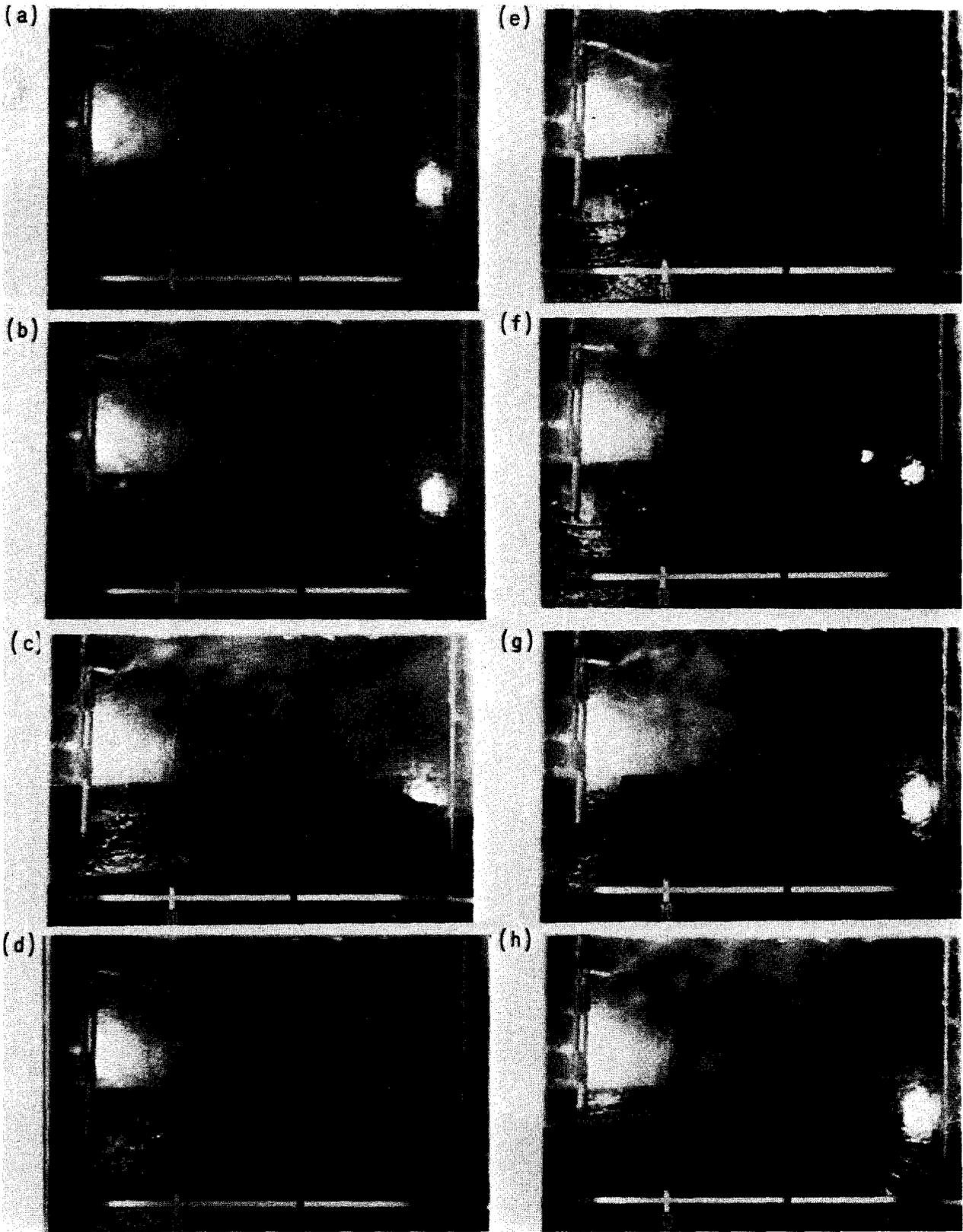


Figure 15.—Smoke tracers over well-developed waves.

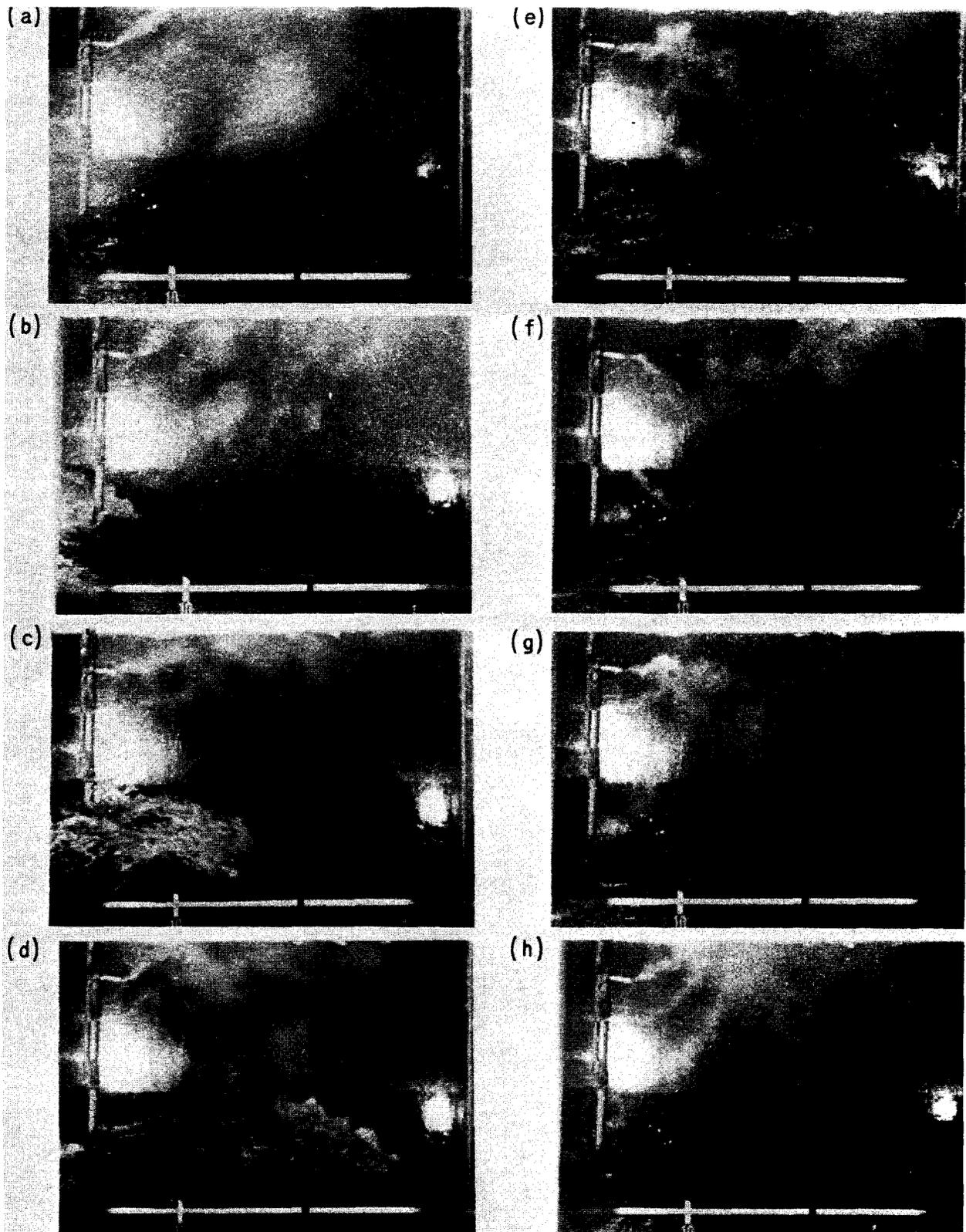


Figure 16.—Smoke tracers over breaking waves.

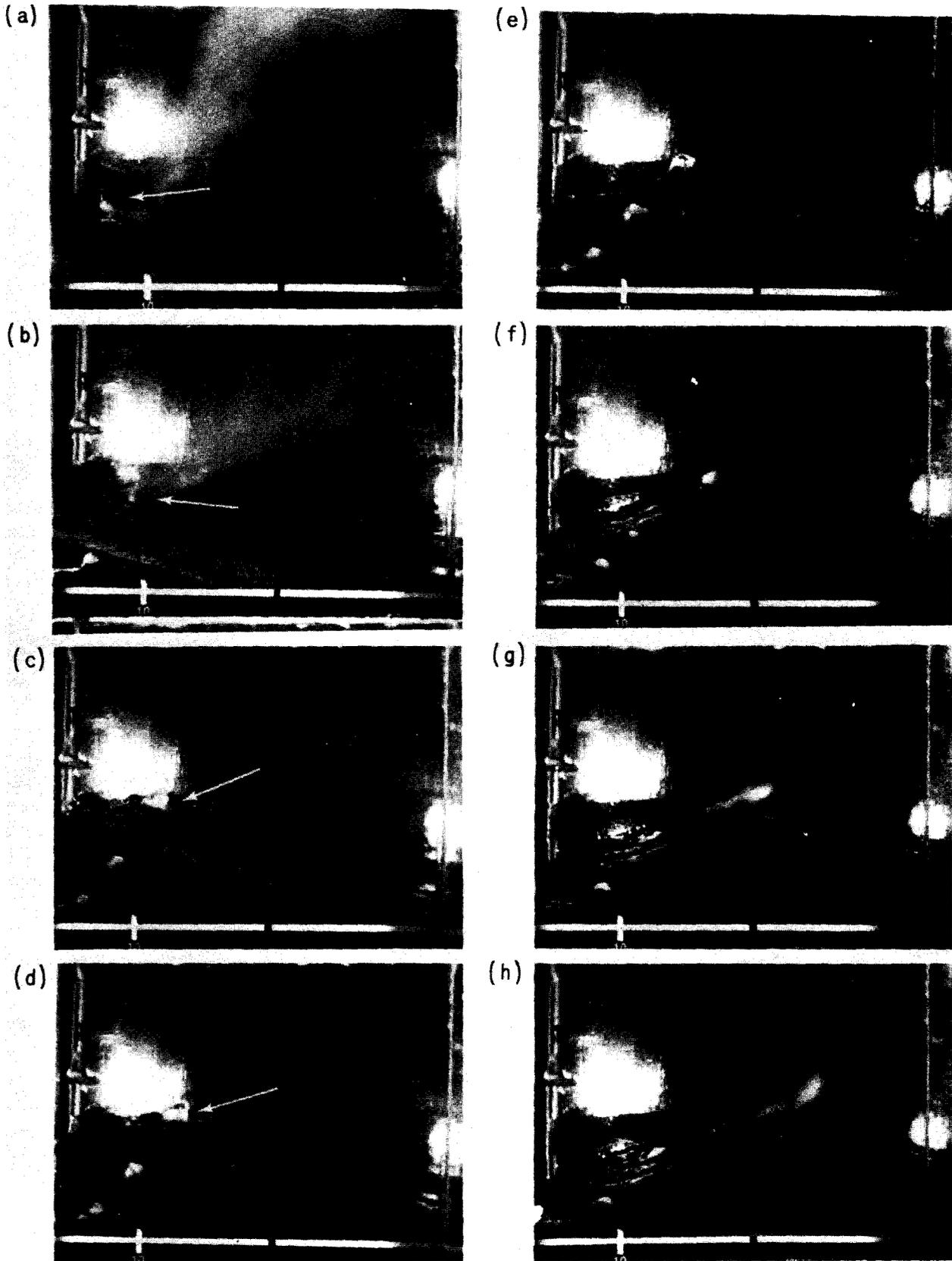


Figure 17.—Smoke bubble released from under water.

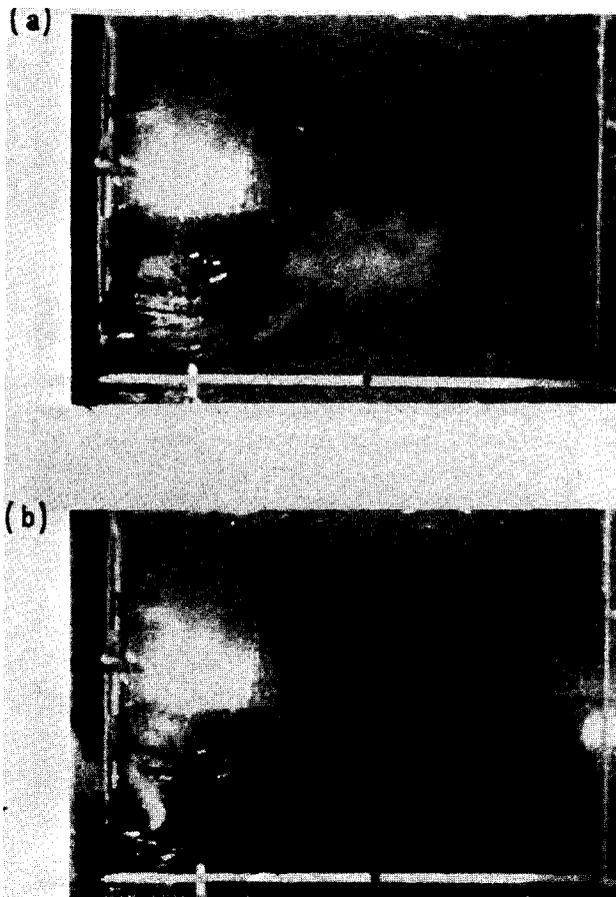


Figure 18.—Smoke tracers showing reverse flow of air near surface ahead of approaching wave crest.

the upwind side of the crest. The trough region is characterized by rather random air motion at this probe level (about 10 inches above the undisturbed water surface). The flow then shifts upward again ahead of the next wave crest (figures 15g–h). Figure 16 presents a similar series of frames but with a heavily breaking wave. The vapor tracer pattern is almost identical to that shown in figure 15.

The vapor dispensing nozzle was then lowered to a point where it would be just above the surface in a wave trough and below the surface during the passage of a wave crest. Figure 17 shows pictures taken during the passage of a wave crest with vapor bubbles being fed from below. Note the small cloud of vapor in figure 17a and that it does not disperse at this point but rides over the crest more or less undisturbed until it reaches the upwind side of the wave, figure 17f. In frames *g* and *h* the vapor is quickly blown away. This apparent dead air region at the surface and to the lee of the wave crest was evident in many of the picture series taken. Figure 18 shows

typical flow near the surface in the trough region. Frame *a* traces the motion in the trough, while frame *b* shows the reverse flow building up as the crest approaches.

We have assembled the information obtained from the several hundred feet of film that were taken of bubble trajectories and vapor tracers and constructed a model of the air flow relative to the wave contour. This model is shown in figure 19. The most noticeable feature of the model is the twin vortex system (*c*), on the downwind slope of the wave. The region marked (*a*) is apparently the only place where the main air flow contacts the surface. Regions (*b*) are “dead air” spots where there is little or no air motion relative to the wave. This model, based on data obtained with gravity waves, is consistent with the flow description above capillary waves reported by Schooley.

It should be emphasized that the model shown in figure 19 is only tentative since the data samples were relatively sparse. There is no way of knowing from existing data whether this structure exists over each wave or, if it does not, what percentage of time it is present. The study was made at only one combination of wind and wave parameters. The wind speed was 10 feet per second and the wave parameter H/T was set at approximately 0.4 foot per second. This combination was chosen in order to place the conditions near the minimum region on the D factor curves shown in figure 9. Because only a single wind-wave condition was used in the study, we cannot say that the same flow structure exists under other conditions. A complete investigation into this particular problem was not possible within the limitations of the present study. A substantial amount of additional work should be done to analyze in detail the flow over a variety of wave characteristics.

The flow tracing work that was attempted was aimed only at uncovering possible evaporation retarding mechanisms and we think this purpose was served successfully. Whenever the tunnel is operated with substantial wind velocity, small capillaries form on the surface, regardless of the presence of large waves. The microscale processes that occur in response to the capillaries must be about the same with or without the large waves. The changes in the macroscale flow pattern with changes in wave state must be responsible for the decrease in evaporation under some wave conditions. It is apparent from figures 13 and 14 that such macroscale changes do occur. It is postulated that as wave height to length ratio increases, the twin vortices shown in the model begin to form. These vortices are probably unstable

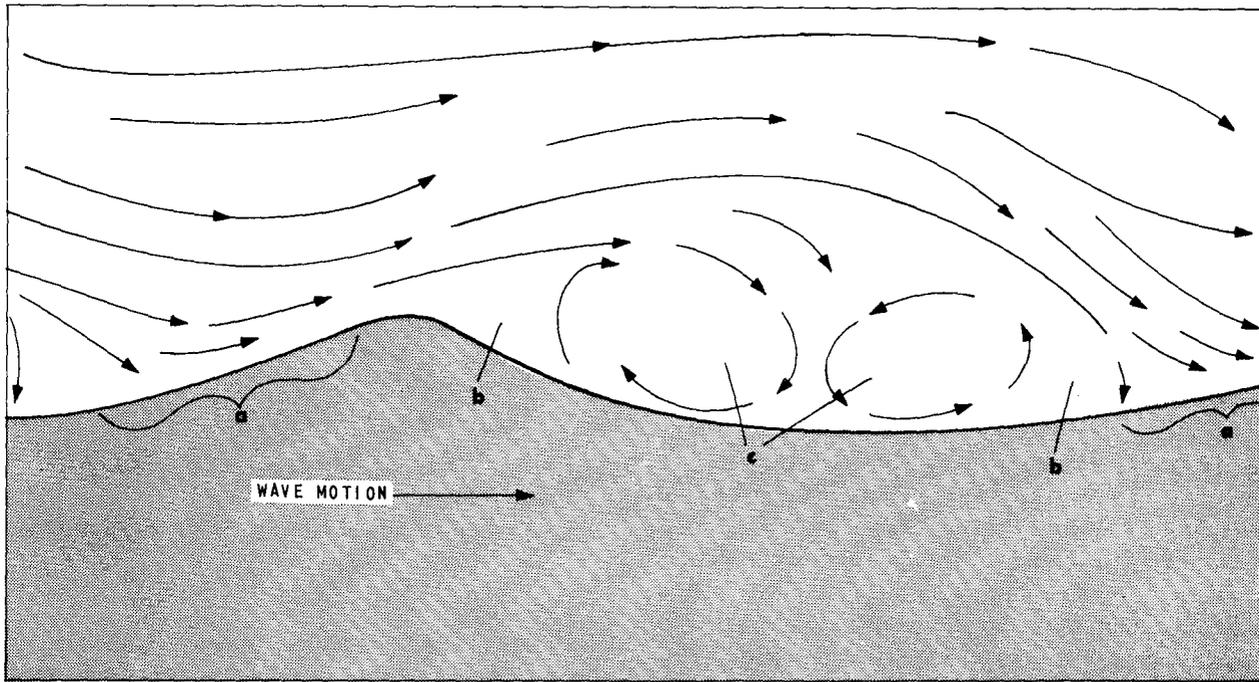


Figure 19.—Model of air flow over well-developed waves.

and break down to cause increased turbulence and evaporation. With further increases in H/L the vortices become more stable so that they limit the transfer of mass between the water surface and the main air flow. This limiting action would be effective over roughly half of the water surface, i.e., on the lee side of the wave crest to about the trough. On the downwind side of the trough the transfer rate need not be significantly affected. As H/L increases further and waves begin to break, the sharp curvature of the water surface due to breaking and bubbling would tend to cause the twin vortices to be shed more easily so that the barrier to mass transfer would be removed and, indeed, more intense turbulence could be generated. Thus the trend would be reversed and the evaporation rate would again increase.

Conclusions

The results of the wave-tank experiments have shown a very definite relationship between surface wave conditions and the measured evaporation rate.

In particular, it appears that certain combinations of wind speed and wave conditions actually lead to evaporation rates somewhat less than those measured with similar wind speed and smaller waves or smaller H/T .

Subsequent investigation of the air flow over the waves has detected the presence of dead air regions and vortices which are apparently trapped in the wave troughs and move along with the wave system. It is thought that the build-up of high humidity in these vortices and adjacent dead air regions tends to reduce the water transfer across the air-water boundary. At the same time, the reduced turbulent mixing between the dead air spaces and the upper flow inhibits vertical transport of moisture.

There is little doubt that the confined nature of the evaporation duct in the laboratory equipment tends to influence the air flow patterns which develop. With this in mind, we cannot simply extend the laboratory results directly to the field situation without supporting measurements. The Lake Hefner evaporation study was undertaken in order to provide the necessary tie-in information.

THE LAKE HEFNER EVAPORATION STUDY

In order to fulfill all objectives of the study as outlined in the Introduction, it was necessary to extend the laboratory work into a full-scale field study. The Bureau of Reclamation had already organized an extensive evaporation study at Lake Hefner, Oklahoma, during the summer of 1966. Several different groups were to take part, including the Radio Meteorology Section of the Tropospheric Telecommunications Laboratory of the Environmental Science Services Administration (ESSA). The ESSA group was scheduled to make evaporation measurements which appeared to ideally suit the needs of the CAL experiment. Thus, the CAL wave program was included in the 1966 study.

The program plan called for CAL to make wave measurements at two of the instrument stations at the lake and to use the meteorological records being taken by ESSA for evaporation computations. The instruments installed by ESSA were designed to make the measurements required for evaporation computations by the eddy flux method. The eddy flux equation for evaporation (Swinbank) is

$$E = \overline{\rho_w W} = \bar{\rho}_w \bar{W} + \overline{\rho_w' W'} \quad (21)$$

where ρ_w = vapor density
 W = vertical wind speed.

The primed symbols denote departures from the mean value. The averaging time must be long in comparison to the longest period of fluctuation. The advective term $\bar{\rho}_w \bar{W}$ is small near the surface where $\bar{W} \approx 0$. Thus, we can equate the vapor flux ($\overline{\rho_w W}$) to the eddy flux ($\overline{\rho_w' W'}$). Instrumentation used for measuring ρ_w and W must respond to all fluctuations which contribute to the flux. However, very small-scale fluctuations will not contribute significantly to the vertical flux if the vapor density gradient is small. An indication of the magnitude of error to be expected in the flux estimate caused by slow response of instruments is given by Deacon. The eddy flux method was appropriate for the CAL study because it gives a local evaporation rate, averaged over a relatively short time interval (10 minutes).

The eddy flux evaporation data were used to compute an eddy diffusion coefficient similar to the

one measured directly in the wave tank experiment. The Lake Hefner coefficients were computed using the equation

$$D = \bar{E} / \bar{\rho} (\bar{q}_s - \bar{q}_z) \quad (22)$$

where \bar{E} is the computed vapor flux, $\overline{\rho_w' W'}$

$\bar{\rho}$ = mean air density

\bar{q}_s = saturation mixing ratio at the mean surface temperature

\bar{q}_z = mean value of mixing ratio at level z .

Coincident with the meteorological measurements, time-height profiles of the lake surface level were recorded at each station. From these records, all required wave information could be obtained. The wave data and calculated values of evaporation coefficient were then used to study the relationship between evaporation and wave state.

Figure 20 shows a map of Lake Hefner. The two instrument stations used for the CAL study were the Mid-lake tower station and the Intake tower station. The Mid-lake station is about two-thirds of a mile from the south shore; the Intake tower station is near the north shore with a fetch of about 2.5 miles for southerly winds. Southerly winds predominate in the area during late summer, and the instrument locations were planned to take advantage of this climatological feature.

Instrumentation at Lake Hefner

The instrumentation required for computation of evaporation rates was installed and operated by the ESSA group. The installations of prime importance to the CAL study were the Mid-lake tower and the Intake tower. The Mid-lake tower station was instrumented at heights of 2, 8, and 16 meters. The absolute temperature measurement was made at 8 meters with temperature difference between 8 meters, the surface and the 2-meter level being sensed by thermocouples. The wind velocity measurements were made at all levels with propeller bivanes. The bivanes measured wind speed, azimuth angle θ , and elevation angle ϕ . A sonic anemometer for measuring the vertical wind component, and a microwave

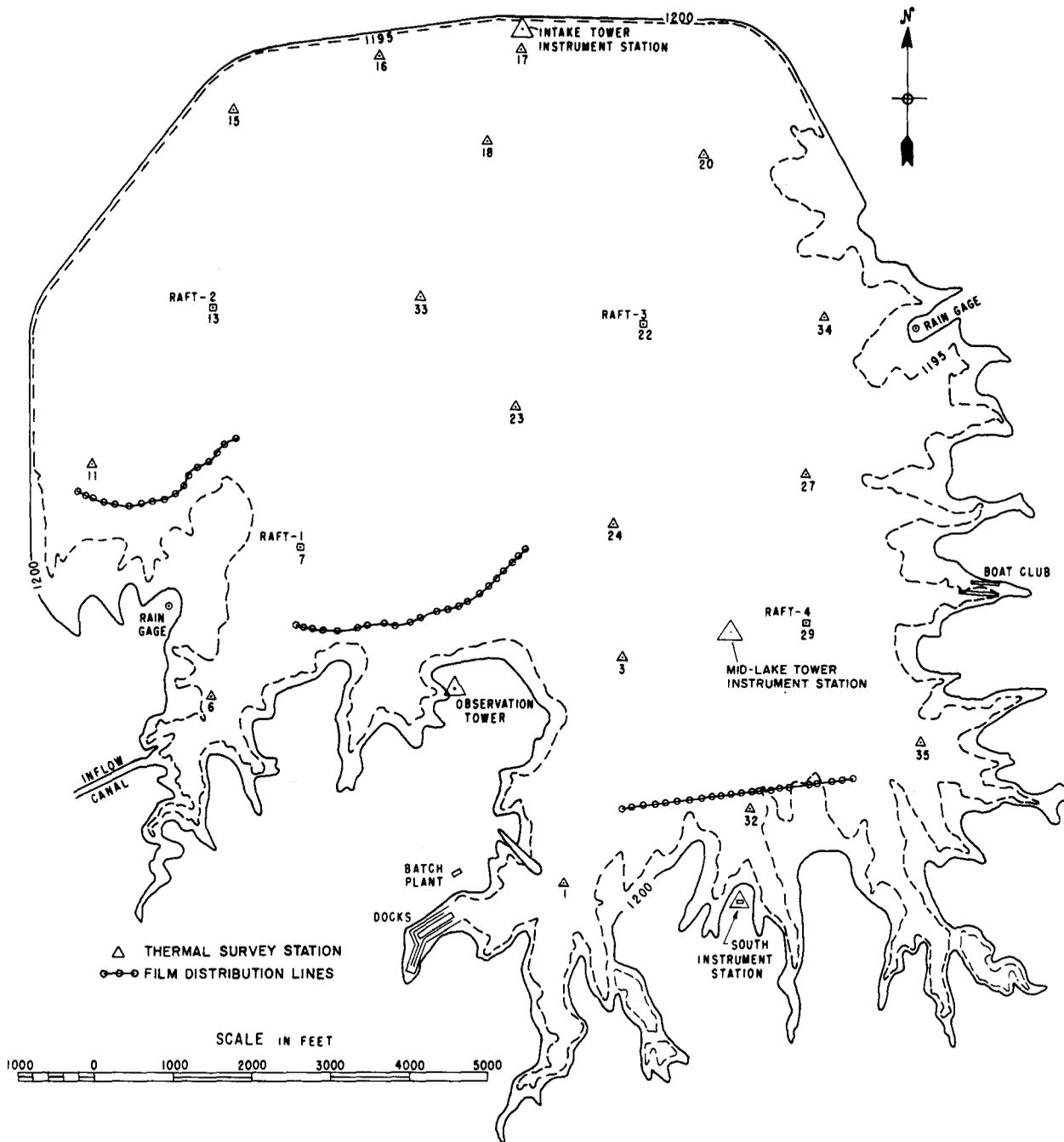


Figure 20.—Map of Lake Hefner, Oklahoma City, Okla.

refractometer for sensing fluctuations in humidity, were located at the 8-meter level. Barium floride resistance strips were used at all levels for measuring the relative humidity.

Similar instruments were installed at the Intake tower with the exception of the sonic anemometer and refractometer. Only the 2-meter and the 8-meter

levels were instrumented at the Intake tower. Measurement stations were also located at the east side and the south side of the lake, but the data from these sites were not used in the CAL study.

Measurement of wave conditions was undertaken by CAL as part of the extended wave-evaporation study. The instrument designed for the wave meas-

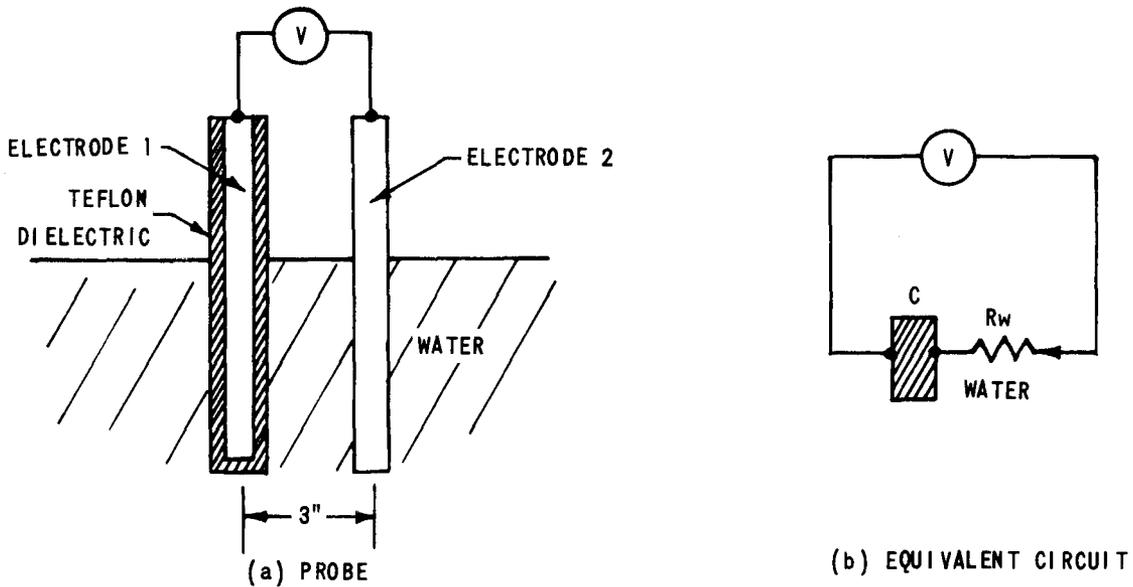


Figure 21.—Basic wave probe configuration.

urement was a capacity-type probe that provides a linear voltage output proportional to the water level on the probe. Recording the output voltage on a time-base record (tape recorder) produces a height-time record of waves, from which power spectra and other wave parameters such as peak-to-trough heights and wave periods can be computed. The probe consists of two long electrodes, one of which is covered with a teflon sheath. Figure 21(a) shows the basic probe configuration.

Electrode 1 is a conducting rod which forms the inner plate of a cylindrical capacitor. Electrode 2 simply makes contact with the water surrounding the outer surface of the teflon dielectric covering, thus making the water the outer plate of the capacitor. Water is not a good conductor (as compared to most metals) so that some resistance is introduced, as shown in the equivalent circuit. Design and construction details for the wave probe are given in appendix C.

Three probes were constructed for the Lake Hefner study. During the 1966 program, one probe was installed at the Mid-lake tower and one at the Intake tower with the third unit being kept as a spare. The wave probe installation at each site is shown in figure 22. Provisions were made in the mounting structure so that the units could be easily raised or lowered with changing water level to keep the mean level at about the mid-point of the probes.

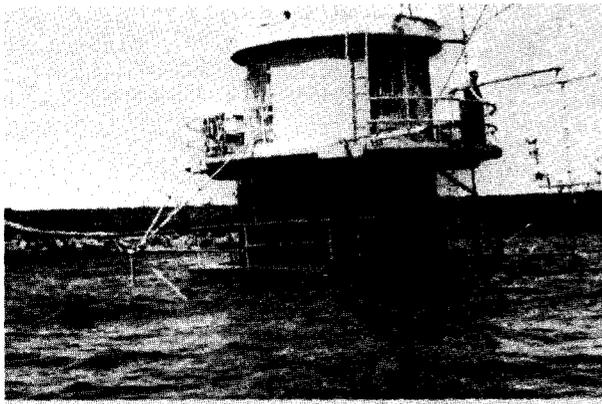
In the August 1967 program, ESSA did not instrument the Intake tower for collection of eddy flux data.

Thus, it was decided not to place wave measuring instrumentation at this point. Since three probes were available, it was decided to install all three in a triangular configuration at the Mid-lake tower. The triangle was formed with one side oriented north-south and another east-west. The multiprobe arrangement was used in order to check wave velocity and possibly direction of motion (parameters that could not be derived from the single probe data taken in 1966). Figure 23 shows the 1967 Mid-lake tower installation. The spacing used on the north-south leg was 4 feet and on the east-west side, 3 feet.

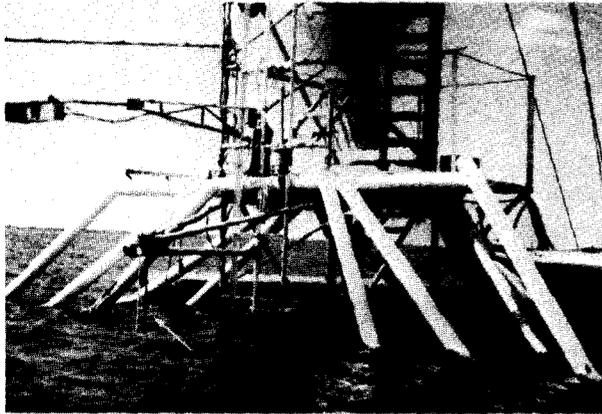
Data Acquisition

In 1966, data recording began on the afternoon of August 25 and ran continuously until a heavy rain forced a shutdown during the early morning of August 31. All instruments appeared to be functioning correctly during the data collection period and over 100 consecutive hours of data were obtained, with only short breaks for calibration checks. An abundance of data was acquired for the medium wind speed range (10 to 20 miles per hour) but very little data were acquired beyond this range. We were especially anxious to obtain data points in the 5 to 10 miles per hour range, and this was the main reason for returning in 1967.

The 1967 data acquisition period began on August 21 and ran intermittently through August 30. Several weather disturbances passed through the area during



(a) INTAKE TOWER



(b) MID-LAKE TOWER

Figure 22.—Wave probe installation at Lake Hefner, 1966.

the period, temporarily shifting the wind out of the south. Nevertheless, several cases with southerly winds in the 5 to 10 miles per hour range were acquired and the CAL observations were terminated at the end of August, although the ESSA group remained with the hope of expanding the data sample.

Processing and Analysis of Lake Hefner Data

During the main data acquisition period in 1966, all data were recorded on ESSA magnetic tape recorders, including the CAL wave probe outputs. For the sake of efficiency and cost it was agreed that CAL would use the data samples selected by ESSA, and that ESSA would compute the power spectra of the wave records and furnish them to us along with the meteorological parameters required to complete the analysis program outlined at the beginning of this section of the report.

In order to extend the data set, CAL personnel

visited ESSA and, using a Precision Instrument tape recorder, directly recorded 22 more samples (11 for each of the two sites).

During the 1967 field program the wave data were recorded on a CAL recorder, along with the time code signal from the ESSA equipment. We then copied the required 1967 meteorological data along with the time code signal from ESSA tapes. The common time code on the two sets of tapes provided the means to time synchronize the two sets of data during the conversion to digital form for machine processing.

1. Initial data processing

The raw data required for the analyses were:

- (a) air temperature at 8 meters
- (b) ΔT , surface to 8 meters
- (c) ΔT , 2 meters to 8 meters
- (d) relative humidity at 2 meters
- (e) relative humidity at 8 meters
- (f) wind speed at 2 meters
- (g) elevation angle of 2-meter bivanne
- (h) wind speed at 8 meters
- (i) elevation angle of 8-meter bivanne
- (j) water surface level (wave probe).

Ten-minute sections were selected from the analog data records and converted to digital form using a sampling rate of 5 per second. This rate provided 3,000 data points per set. The digital data were then processed by computer. The outputs of the computer program were:

- (a) mean surface temperature, \bar{T}_s
- (b) mean temperature at 2 meters, \bar{T}_2
- (c) mean vapor density at 2 meters, $\bar{\rho}_{w2}$
- (d) mean wind speed at 2 meters, \bar{V}_2
- (e) mean vertical eddy flux of water vapor at 2 meters, $\overline{\rho_{w2}' W_2'}$
- (f) mean temperature at 8 meters, \bar{T}_8
- (g) mean vapor density at 8 meters, $\bar{\rho}_{w8}$
- (h) mean wind speed at 8 meters, \bar{V}_8
- (i) mean vertical flux of water vapor at 8 meters, $\overline{\rho_{w8}' W_8'}$
- (j) power spectrum of the water surface record $[P_{WH}(f)]$.

The wave power spectra were computed using the equal-time-spaced, discrete sample method of Blackman and Tukey. In addition to the wave power spectrum, the variance of the wave record (σ^2_{WH}) was also computed. By definition

$$\sigma^2_{WH} = \int_{-\infty}^{+\infty} P_{WH}(f) df \quad (23)$$

and represents the total energy in the spectrum.

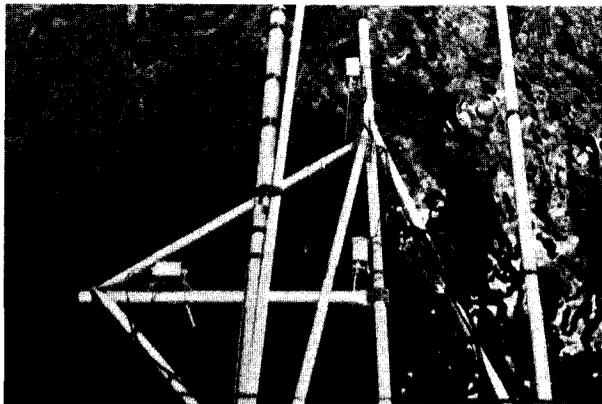
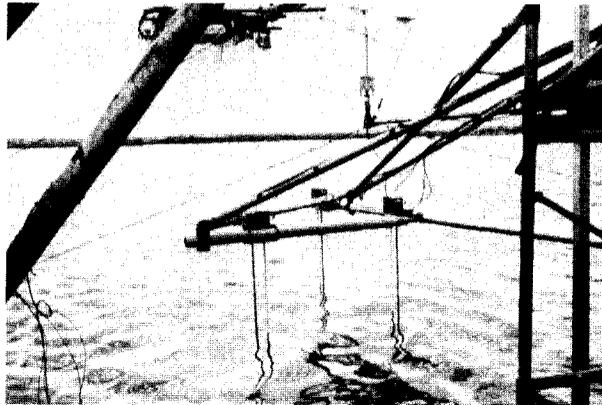
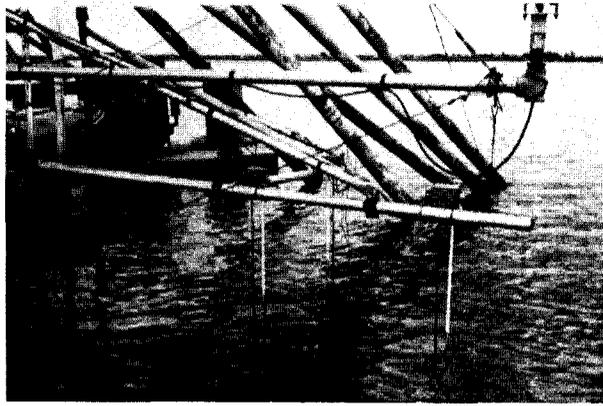


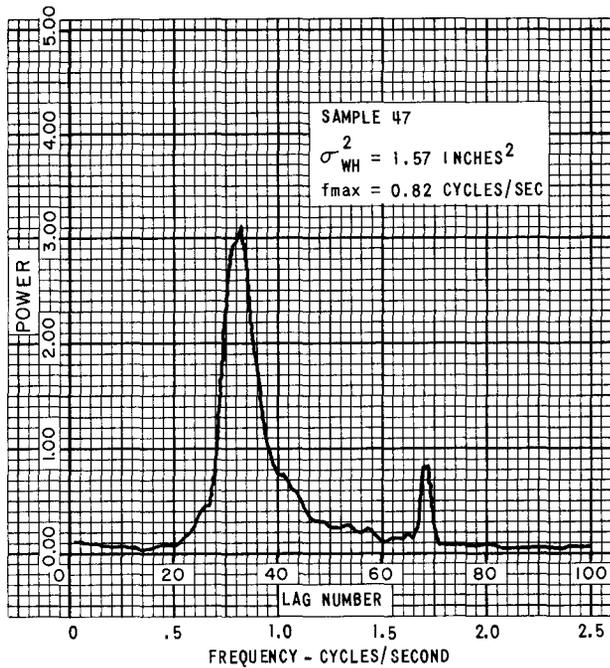
Figure 23.—Wave probe installation at the Mid-lake tower, 1967.

Figure 24 shows sample spectra computed from the Lake Hefner data. The numbers plotted along the abscissa are lag numbers from the computation process and represent frequency. The frequency corresponding to the lag number can be computed from

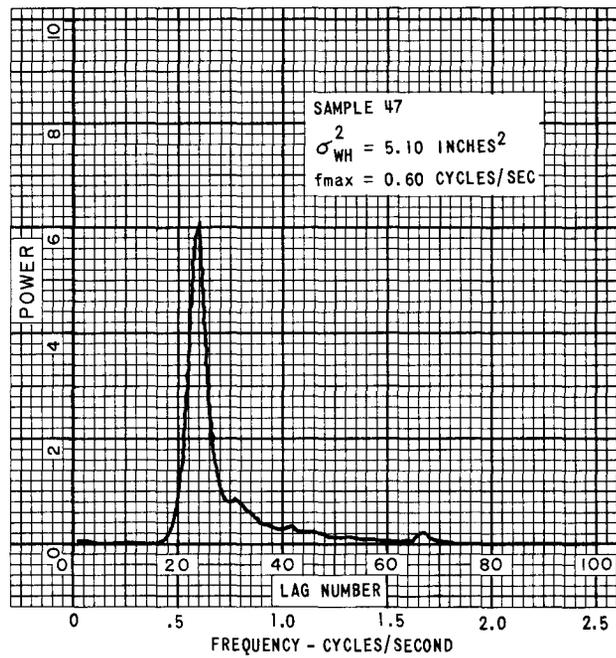
the simple relationship,

$$\text{frequency} = \frac{\text{lag number}}{100} \times 2.5 \text{ cycles per second}$$

The spectra have all been normalized to unit area



a) Mid-Lake Tower



b) Intake Tower

Figure 24.—Sample wave spectra, August 1966.

by dividing each spectral estimate by the variance, e.g.,

$$\frac{1}{\sigma^2_{WH}} \int_{-\infty}^{+\infty} P_{WH}(f) df = 1 \quad (24)$$

The spectra were plotted by computer, the normalization being programmed in order to better display the plotted curves which cover a very wide range of energies. The scale on the ordinate is seconds per cycle. In order to reintroduce the absolute scale on the ordinate, one needs only to multiply by the variance. The two spectra shown in figure 24 are for the same time period, (a) being taken at the Mid-lake tower, and (b) at the Intake tower. Note that the total wave power increased by a factor of 3.25 between the two sites, and the frequency of the peak decreased from 0.81 cycle per second to 0.6 cycle per second.

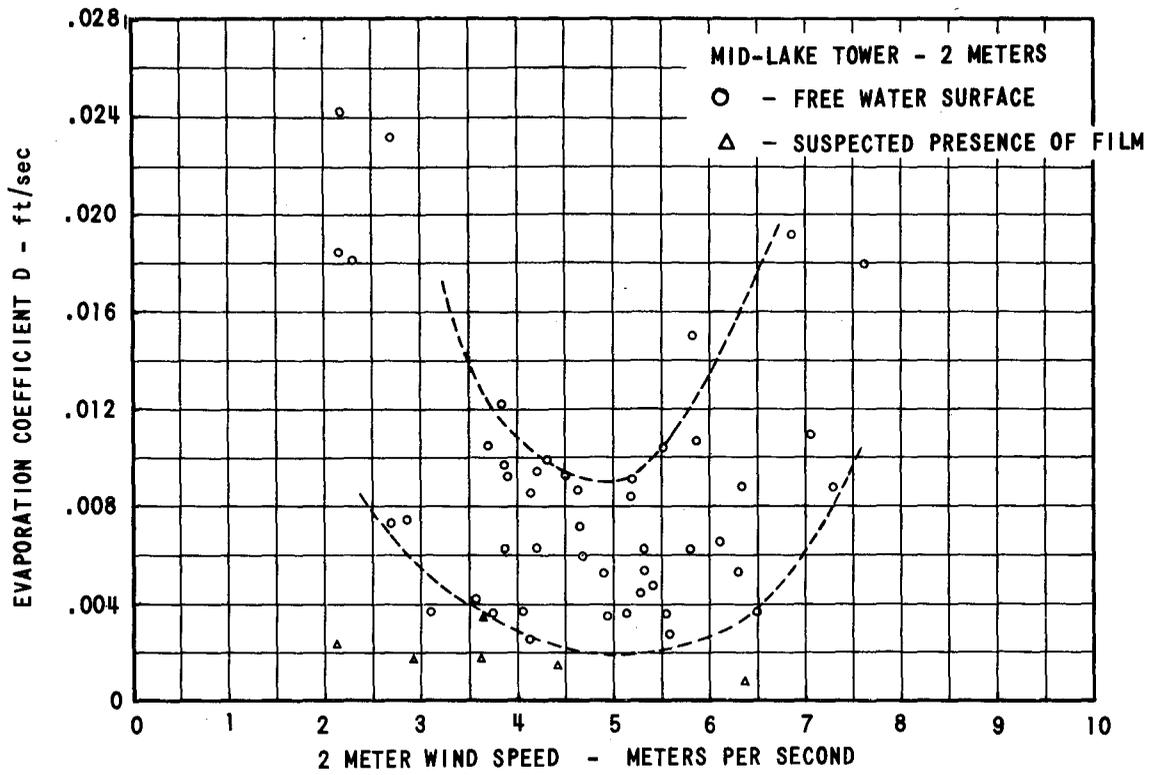
The remainder of the data processing was done by hand. The parameters required to compute D from equation (22) were derived from the computer output data listed above. The value of \bar{q}_s was obtained from tables using \bar{T}_s and a correction for pressure, and \bar{q}_z was computed from the mean vapor density ρ_w . The evaporation rate \bar{E} was equated directly to the vertical flux of water vapor $\rho_w' W'$. All processed data have been tabulated and form appendix E of this report.

2. Analysis of the processed Lake Hefner data

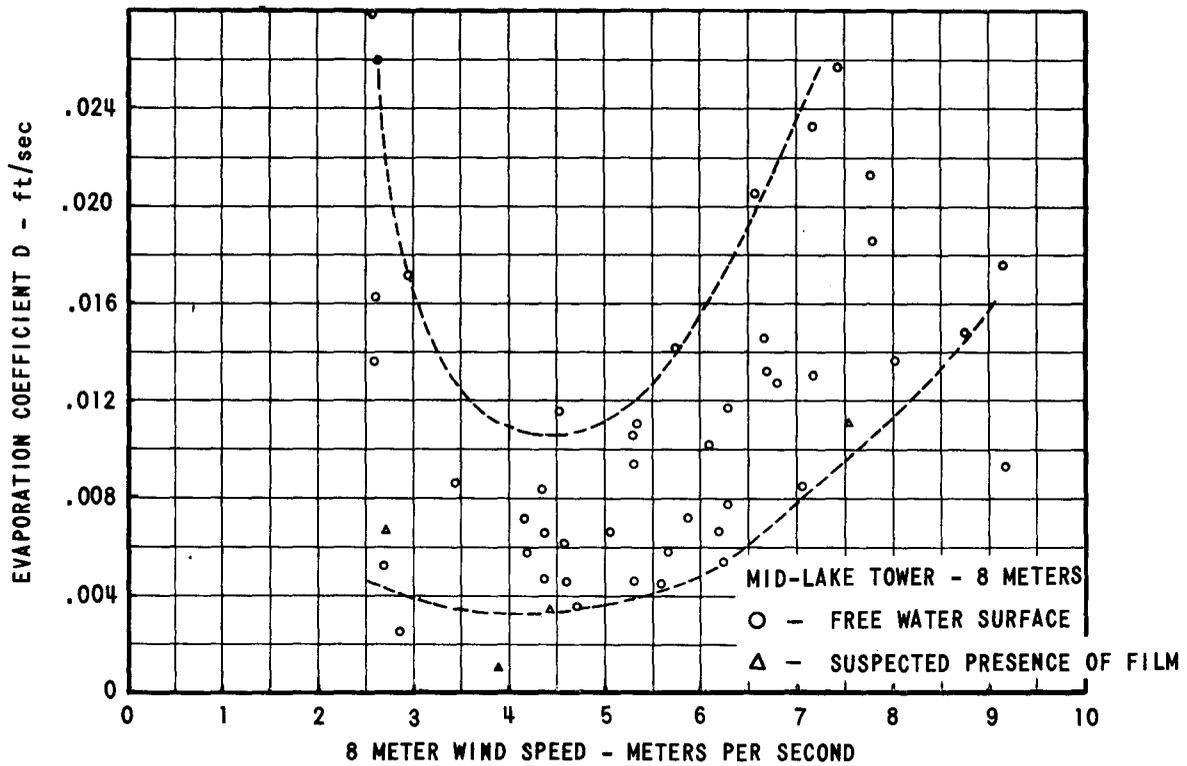
(a) Evaporation coefficient and wind speed

One of the prime objectives of the Lake Hefner work was to check the validity of the wave tank results in a full-scale environment free from restricting boundaries and artificially generated ambient conditions. The first check that can be made is simply to plot the computed coefficient D , as defined by equation (22), as a function of wind speed. Previous results from tests of mass transport equations indicate that D , as defined, should be a linear function of wind speed (Deacon and Webb). However, if the wave tank data have any meaning in the "real world" we would predict a possible departure from linearity in a field situation due to the wave effect. (The wave parameters at a fixed point in the lake under equilibrium conditions are necessarily some function of the wind velocity.)

Figures 25 and 26 are plots of D as a function of V , constructed from measurements taken at 2 and 8 meters for both the Mid-lake tower and the Intake tower. The rather large amount of scatter in the data is certainly disturbing. Nevertheless, non-linearity in each of the four data envelopes is fairly obvious. Note that the minimum in the data for the Intake tower at 2 meters is especially sharp. While it could be argued that the Intake tower data might fit a straight line, the Mid-lake tower data leave no

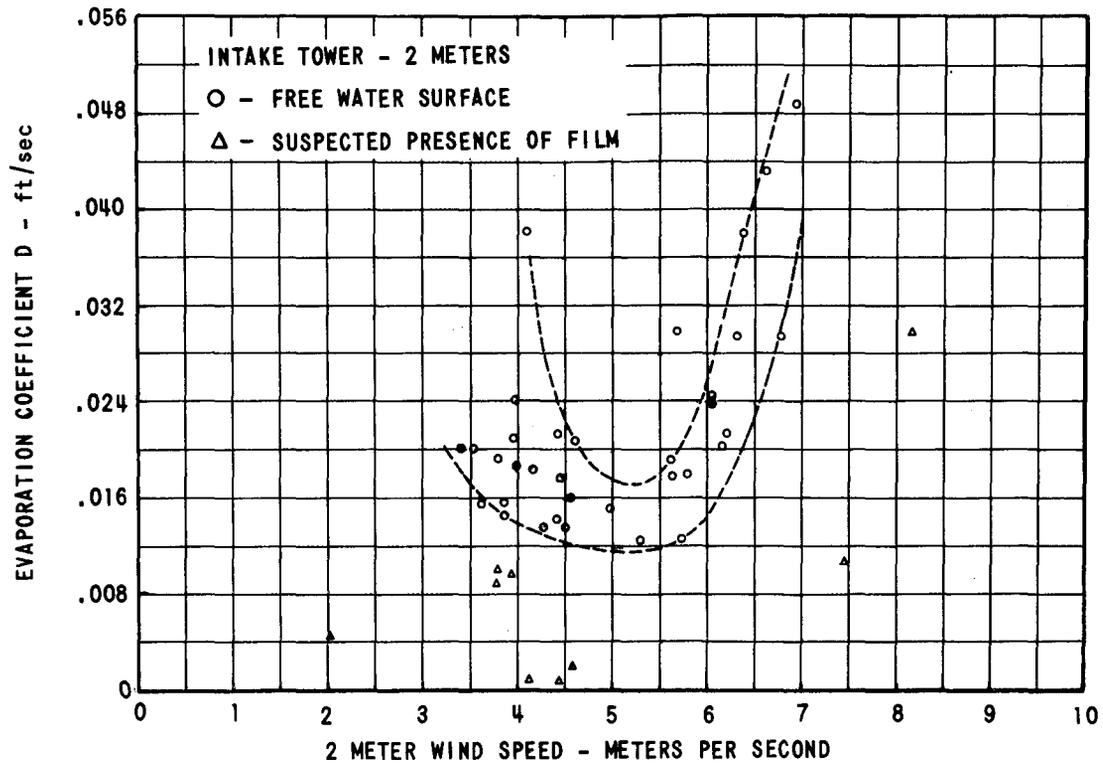


a) Computed From The 2 Meter Level Data

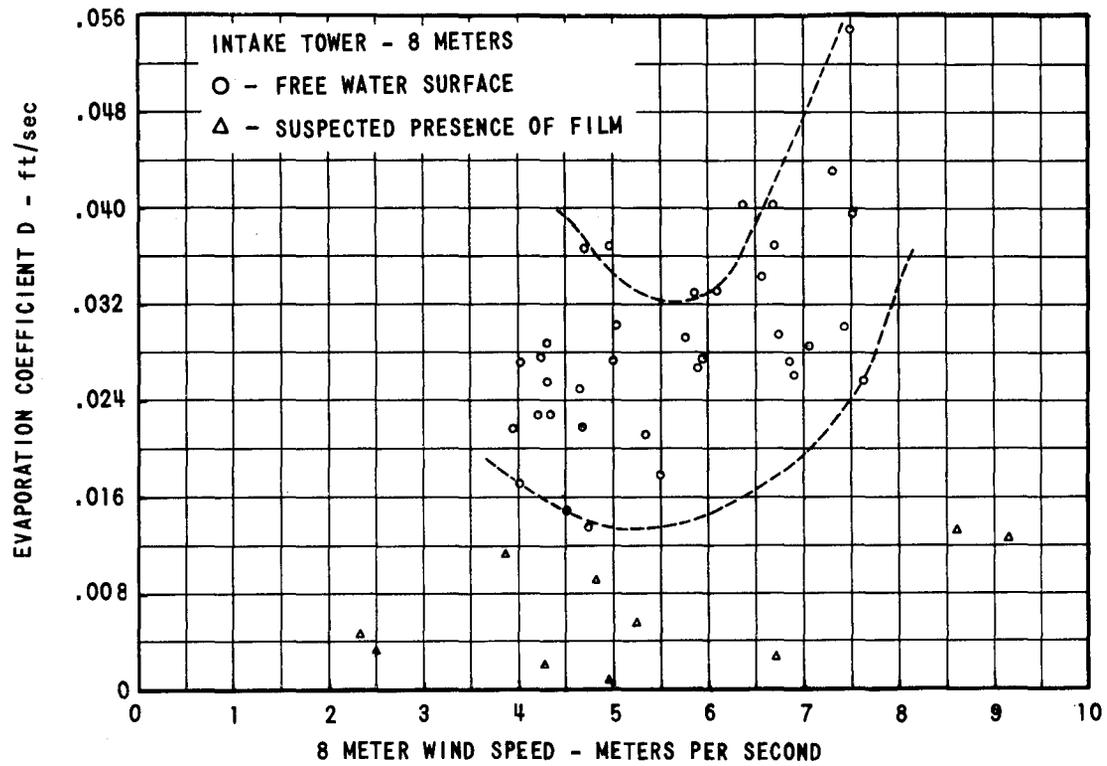


b) Computed From The 8 Meter Level Data

Figure 25.—Evaporation coefficient D as a function of wind speed, Mid-lake tower.



a) Computed From The 2 Meter Level Data



b) Computed From The 8 Meter Level Data

Figure 26.—Evaporation coefficient D as a function of wind speed, Intake tower.

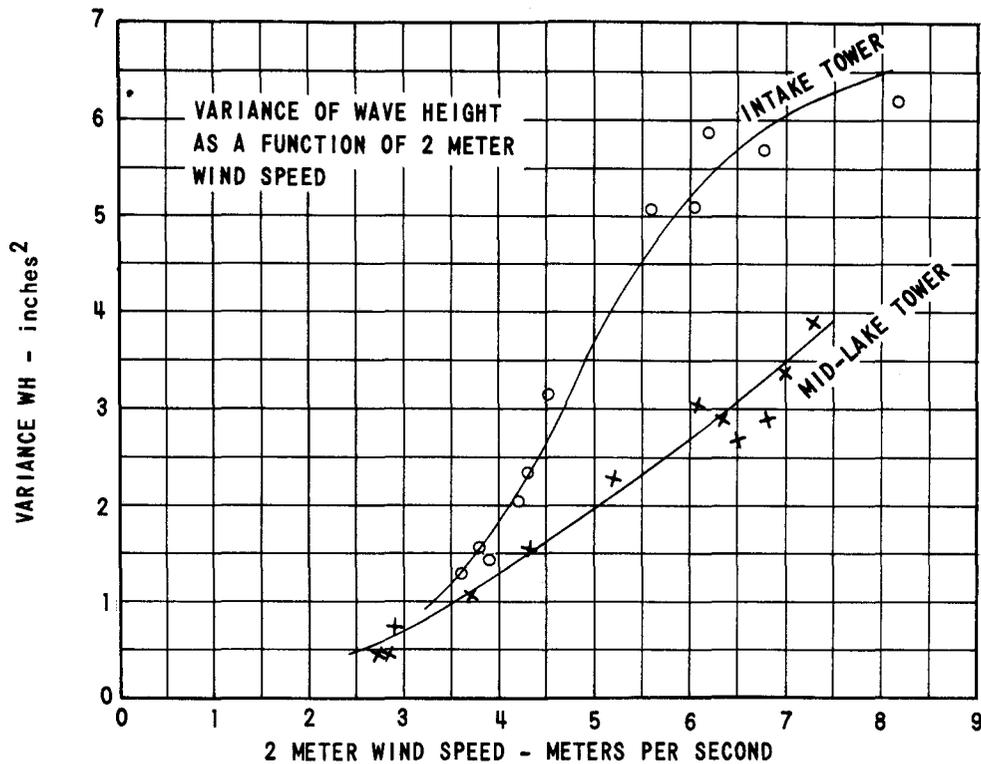


Figure 27.—Total wave energy as a function of the 2-meter wind speed.

doubt about the existence of a minimum in D for wind speeds of 4 to 5 meters per second.

(b) Wave data analysis

In order to make a comparison between laboratory results and Lake Hefner data, the Lake Hefner wind-wave relationships are required. Twenty-four wave power spectra were computed, representing the full range of wind speeds encountered during the field studies. The spectra were all strongly peaked with little wave energy outside the significant wave frequency. The peak frequency could be read very easily from each spectrum. This frequency was determined and converted to wave period.

Along with wave spectra, the variance or total wave energy was also computed. By taking the square root of the variance we obtain the standard deviation of the wave record. Since the spectra indicate very strongly monochromatic waves, and since the waves are nearly sinusoidal in shape, it is not unrealistic to equate the measured standard deviation to the rms value of an equivalent sine wave. The peak-to-trough wave height may then be estimated using the expression:

$$H = 2\sqrt{2}\sigma_{WH} \quad (25)$$

The wave spectra and the variance thus provided the necessary information for computing the wave parameter H/T . Figure 27 presents a plot of wave energy (σ_{WH}^2) as a function of 2-meter wind speed for both the Mid-lake tower and the Intake tower. Note that at a wind speed of 7 to 8 meters per second the wave energy at the Mid-lake tower is still increasing at a near constant rate, while the wave energy at the Intake tower is beginning to reach a maximum. Figure 28 shows the wave parameters or wave height-to-period ratio (H/T) for both sites. Again the Mid-lake tower curve is linear, while that for the Intake tower has a parabolic shape due to the amplitude limiting occurring over the much longer fetch.

Since the wave parameter versus wind speed curve is linear for the Mid-lake tower, a plot of the coefficient D versus H/T would show a distribution similar to those in figures 25 and 26. However, this is not necessarily true for the Intake tower. Figure 29 shows a plot of D as a function of H/T for the 2-meter level at the Intake tower. Although the (H/T , V) relationship is nonlinear, this has not appreciably altered the D coefficient envelope. The gross departure from linearity and the minimum (at $H/T=0.3$) remains as in the wind-dependent curves.

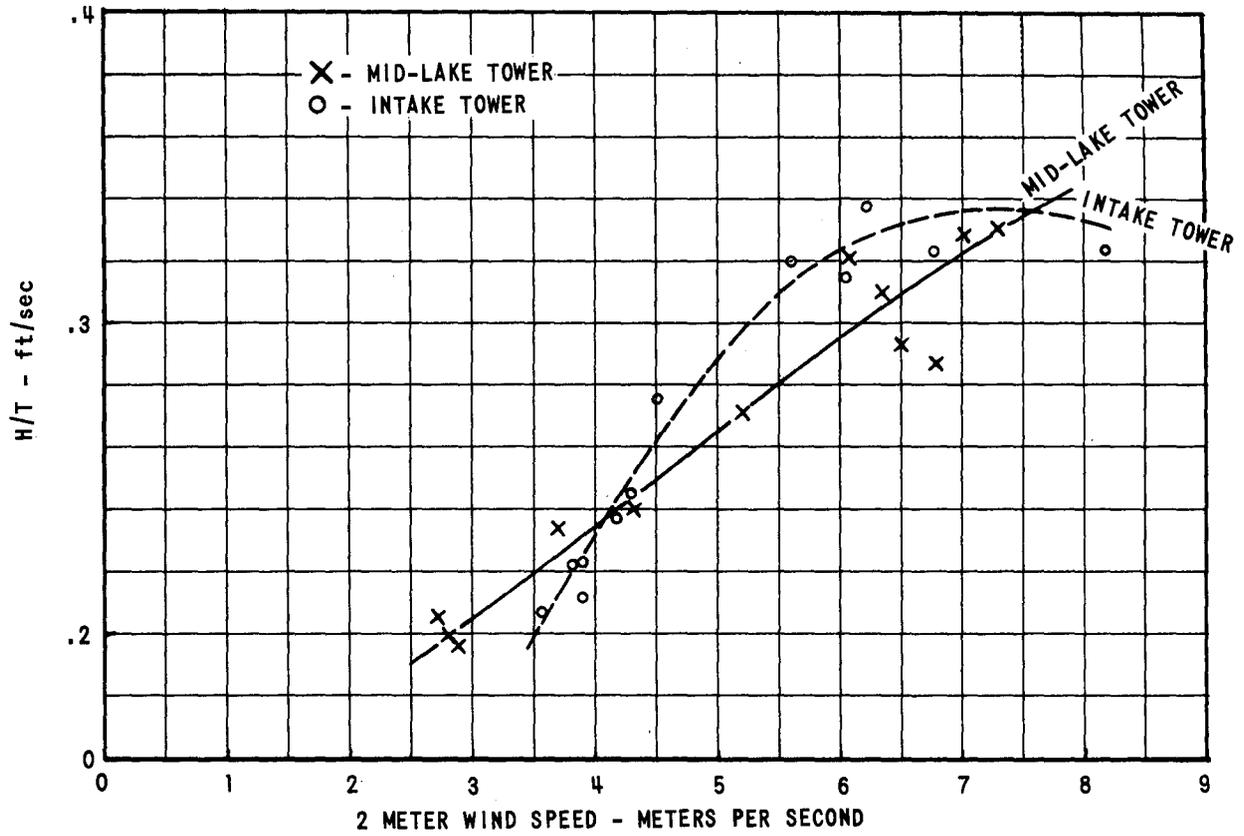


Figure 28.—Wave parameter H/T as a function of the 2-meter wind speed.

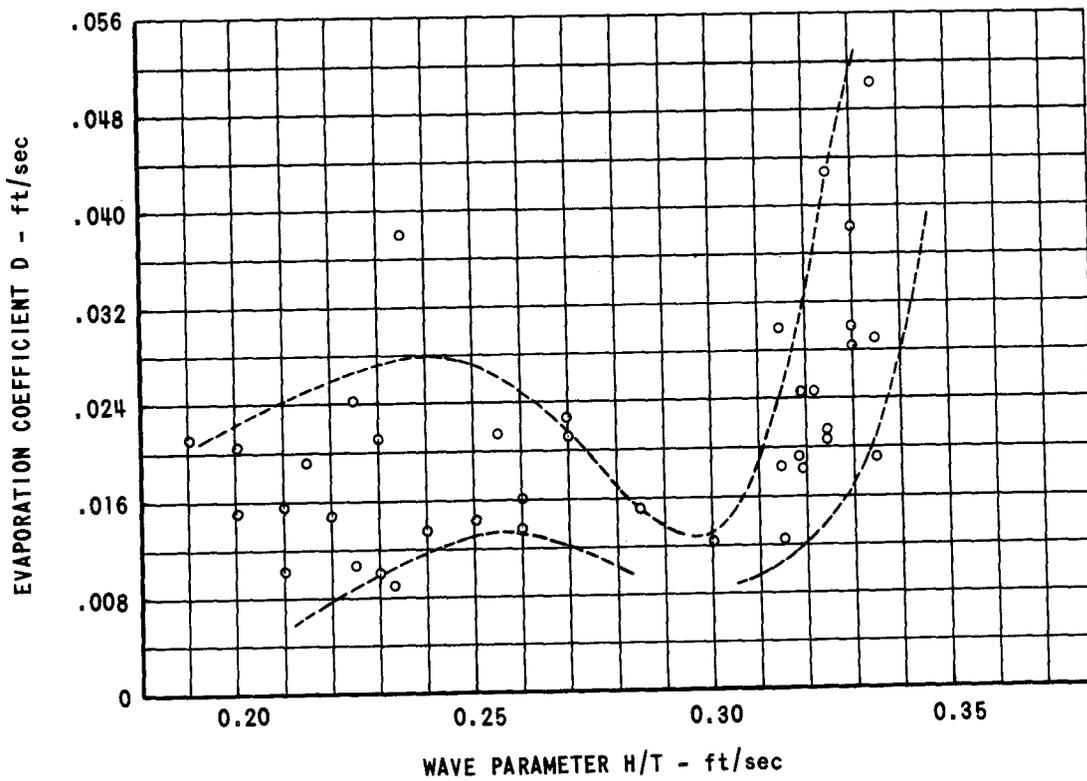


Figure 29.—Evaporation coefficient D as a function of the wave parameter H/T for the 2-meter level at the Intake tower.

COMPARISON OF THE LABORATORY AND FIELD RESULTS

In the laboratory experiments, there was little correlation between wind and waves since they were generated independently. Because there was no way of knowing just what range of wave characteristics would correspond to given wind velocities in the field, the tank experiments were run over the complete range of wind speed and wave parameters attainable with the apparatus. Many of these data lie outside the range likely to occur in the field. We can, however, compare the two sets of data through the wave measurements made at Lake Hefner.

The H/T versus V curves for Lake Hefner (figure 28) were plotted on the $D(H/T, V)$ field computed from the wave tank measurements (figure 9). This combination makes possible the construction of a D versus wind speed curve using the tank evaporation data. As was anticipated, this curve showed a substantially different relationship from that found on Lake Hefner (figures 25 and 26). However, it is pos-

sible to reconcile the two sets of data by a relatively simple adjustment of the wave tank results. Figures 30(a) and 30(b) show the $D(H/T, V)$ field required to duplicate the Lake Hefner results. Note that the general shape of the D contours remains the same as in the wave tank; it has been necessary only to rotate the field clockwise about the origin and scale the coefficient values by a fixed amount. The reason for the scaling factor is quite understandable and was discussed previously under "Results of the Wave Tank Evaporation Experiments." On the other hand, the necessity to rotate the wave-tank-derived D contours is most surprising. As is readily seen in figures 30(a) and 30(b), the rotation process has increased the negative slope to the left of the peak. Thus, we are led to conclude that in the field situation, the influence on evaporation by waves on the surface appears to be even greater than in the wave tank.

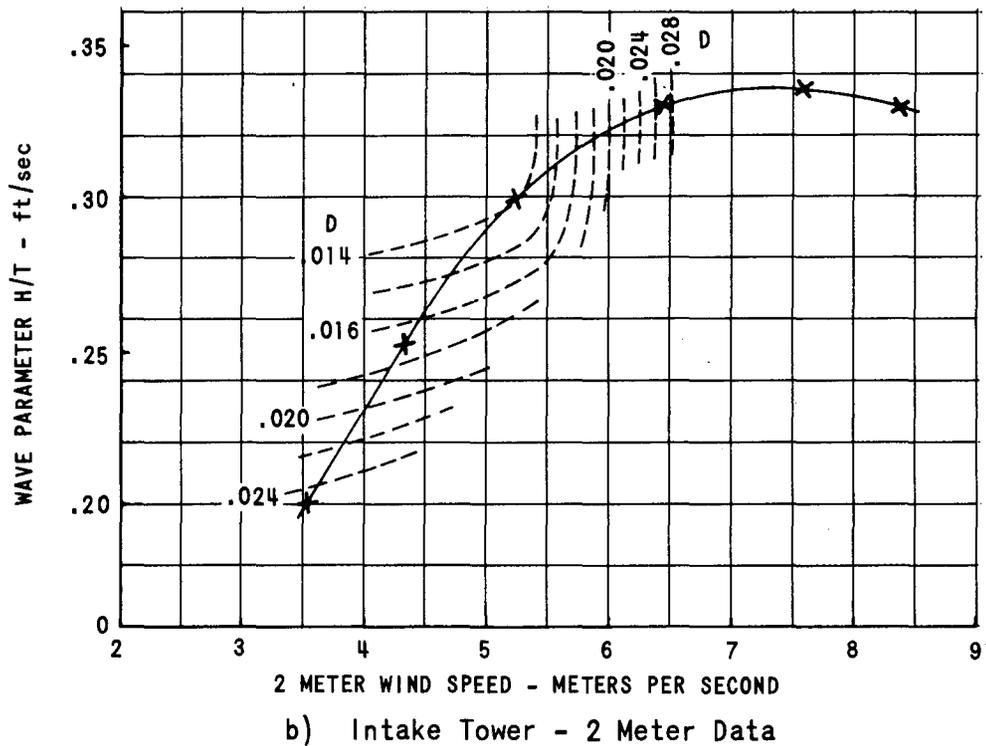
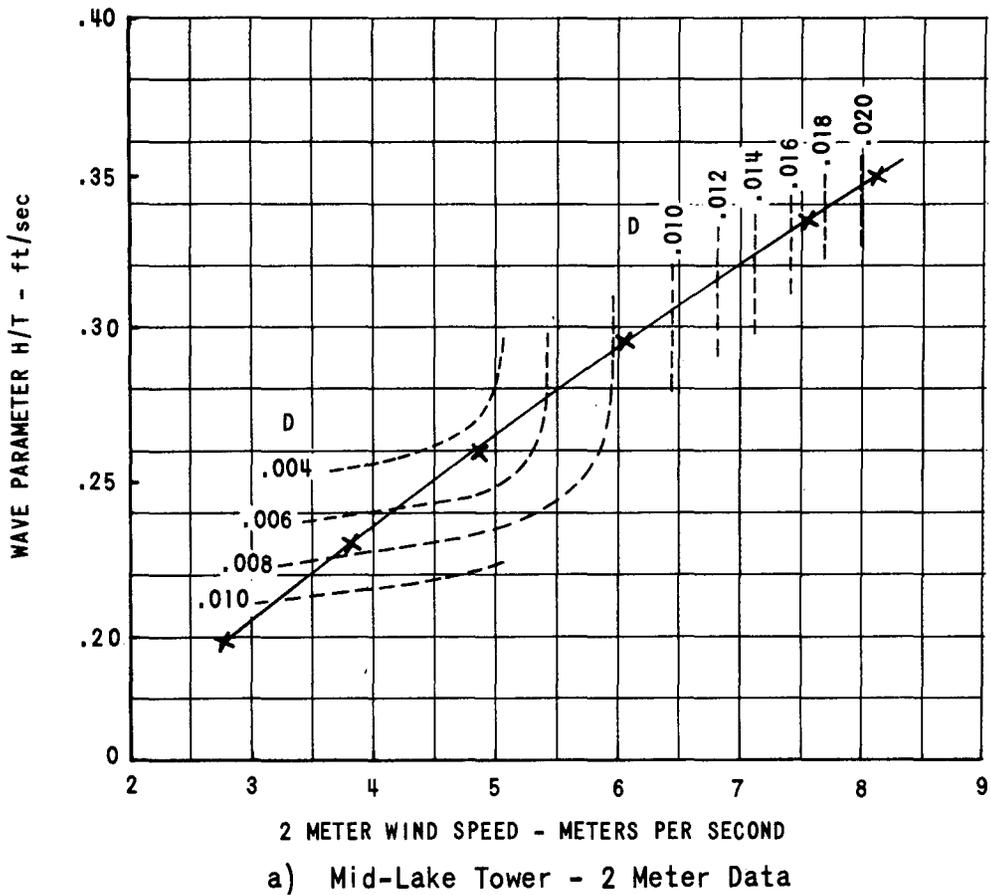


Figure 30.—Contours of constant D as a function of wave parameter H/T and wind speed—contours adjusted to fit the Lake Hefner results.

EVAPORATION ESTIMATES AND WAVES

Although every attempt was made to make the Lake Hefner study as quantitative as possible, there was considerable scatter in the processed data and the overall evaporation estimates were somewhat lower than existing water budget information would have predicted. The reasons for the apparent inaccuracies in the evaporation estimates are very likely due to departure from the ideal conditions assumed in developing the eddy flux theory. For example, neglecting the advective term $\bar{p}_w \bar{W}$ by assuming $\bar{W}=0$, does not necessarily hold. In fact, the mean vertical wind was never zero in any of the processed samples, although this could have been due to small errors in calibration of the bivanes. There was also some evidence that the humidity and wind measuring instruments did not have fast enough response times to follow all the significant eddy scales. Comparison of the measured vapor fluxes from the Mid-lake tower and the Intake tower tend to suggest that the Mid-lake instruments were located on the outer fringe of the vapor blanket. Significantly lower values of the vapor density variance at the Mid-lake tower leads us to this conclusion. The apparent thinness of the vapor blanket at the Mid-lake tower (if this is truly the case) may be due to the reduction in vertical transport of water vapor caused by the waves on the water surface.

While the absolute measurements of evaporation may be somewhat in error, changes in the measured parameters relative to some chosen base value are thought to be more reliable. Thus, we have chosen to use the shape of the measured curve of mass transfer coefficient as a function of wind speed (and the dependent wave parameters), and to adjust the absolute magnitude using the many available water-budget-derived values.

Marciano and Harbeck derived two empirical equations of best fit for Lake Hefner, using slightly different data for each. These are,

$$E = 6.25 \times 10^{-4} V_8 (e_0 - e_8) \quad (26)$$

and

$$E = 6.47 \times 10^{-4} V_a (e_0 - e_a) \quad (27)$$

where E = evaporation in centimeters per 3 hours
 V_8 = wind speed in knots at 8 meters
 Δe = vapor pressure difference in millibars.

Linsley, Kohler, and Paulhus also derived two equations from slightly different data inputs,

$$E = 0.0034 (e_s - e_2) V_4 \quad (28)$$

and

$$E = 0.00270 (e_s - e_2) V_4 \quad (29)$$

where E = evaporation in inches per day
 V_4 = wind speed in miles per day at 4 meters
 Δe = vapor pressure difference in inches of mercury.

During 1965, Oklahoma State University, working for the Bureau of Reclamation, did yet another survey of Lake Hefner. Fry reports two more estimates of the mass transport coefficients, using both water budget and energy budget methods. The water budget form is:

$$E = 12.91 \times 10^{-5} V_2 (e_s - e_2) \quad (30)$$

where E = evaporation in centimeters per day
 V_2 = wind speed in kilometers per day at 2 meters
 Δe = vapor pressure difference in millibars.

All the above equations report a different value of N because they each use different units. Since none of the above forms match the units used in the CAL study, we have chosen to transform these equations to fit our units. The equations reduce to the following (in the order presented):

$$E = 5.38 \times 10^{-5} V_8 (q_s - q_8)$$

$$E = 5.56 \times 10^{-5} V_a (q_s - q_a)$$

$$E = 6.65 \times 10^{-5} V_4 (q_s - q_2)$$

$$E = 5.92 \times 10^{-5} V_4 (q_s - q_2)$$

$$E = 6.17 \times 10^{-5} V_2 (q_s - q_2)$$

where the units are E : grams per centimeter squared per second = centimeters decrease in water level per second

V : feet per second

q : mixing ratio in grams per gram.

Note the close agreement of the equations in spite of the slightly different measuring heights. Normally we would expect the term $V\Delta q$ to be somewhat larger for 8 meters than for 2 meters, leading to lower values of N using 8-meter data. Although there is evidence of this effect in the above equations, the total spread in the values is small. In fact, the total range of values does not exceed the probable error in any individual estimate. Therefore, we have chosen to combine the estimates as they stand and use the mean value of N as the absolute reference for the CAL data.

The mean coefficients derived from the CAL study which correspond to the above are:

$$\text{Mid-lake tower } E = 3.02 \times 10^{-5} V_2 (q_s - q_2) \quad (31)$$

$$\text{Intake tower } E = 5.43 \times 10^{-5} V_2 (q_s - q_2) \quad (32)$$

Equation (31) was derived from the Mid-lake tower data and equation (32) from the Intake tower measurements. Although the Intake tower coefficient agrees very well with the water budget values given above, the Mid-lake tower measurement appears to be too low. Nevertheless, both measurements agree well enough with the water budget data to justify our acceptance of the quantitative estimate of the changes in the mass transfer coefficient with wind and waves.

The mass transport coefficients (N) were computed for each of the available 10-minute samples from the 1966-67 field programs. These data were plotted against wind speed and average curves determined by inspection, giving more weight to the points which lay closer to the centroid of the group. These plots are shown in figures 31a and 31b. The next step was to combine the average curves from the two sites to obtain an average curve for the whole lake. The resulting curve, while not being absolutely correct in magnitude, is believed to represent the correct variation of N with wind speed and associated wave structure.

The average curve for N as determined by combining the data from the two Lake Hefner stations must now be adjusted to correspond to the mean value of the mass transport coefficient derived from

the water budget studies. The mean value of N for the water budget analyses described previously is $N = 5.93 \times 10^{-5}$ centimeter/foot. In order to make the appropriate adjustments, it is required to know the wind speed distribution for the survey periods from which the water budget values were derived. These distributions were not available. However, we have used the 2-meter wind speed distribution from the present Lake Hefner study as a realistic approximation.

The method of adjustment is as follows:

Let

$$N(V_2) = K + n(V_2) \quad (33)$$

where $n(V_2)$ = the average curve from the data

$N(V_2)$ = required $N(V)$ curve

K = adjustment constant.

It is required to find K such that

$$\overline{N(V_2)} = \int_0^{V_{\max}} P(V_2) N(V_2) dV = 5.93 \times 10^{-5} \quad (34)$$

$P(V_2)$ = 2-meter wind speed probability density.

The averaging is done over the wind speed probability density curve derived from the Lake Hefner wind data.

Substituting equation (33) into equation (34)

$$5.93 \times 10^{-5} = K \int_0^{V_{\max}} P(V_2) dV + \int_0^{V_{\max}} P(V_2) n(V_2) dV \quad (35)$$

but

$$\int_0^{V_{\max}} P(V_2) dV \equiv 1 \quad (36)$$

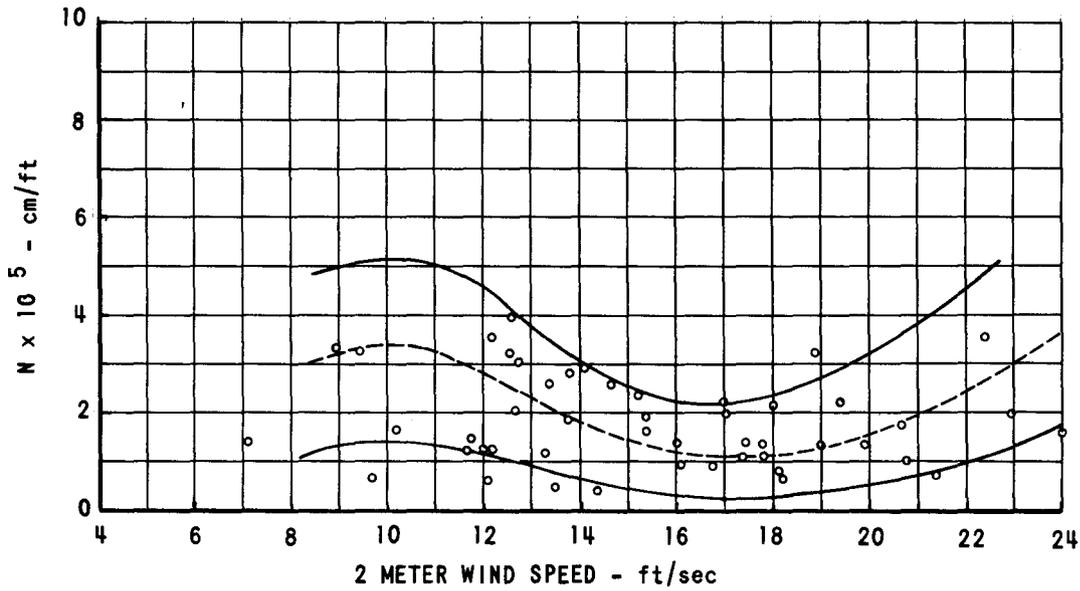
then

$$K = 5.93 \times 10^{-5} - \int_0^{V_{\max}} P(V_2) n(V_2) dV \quad (37)$$

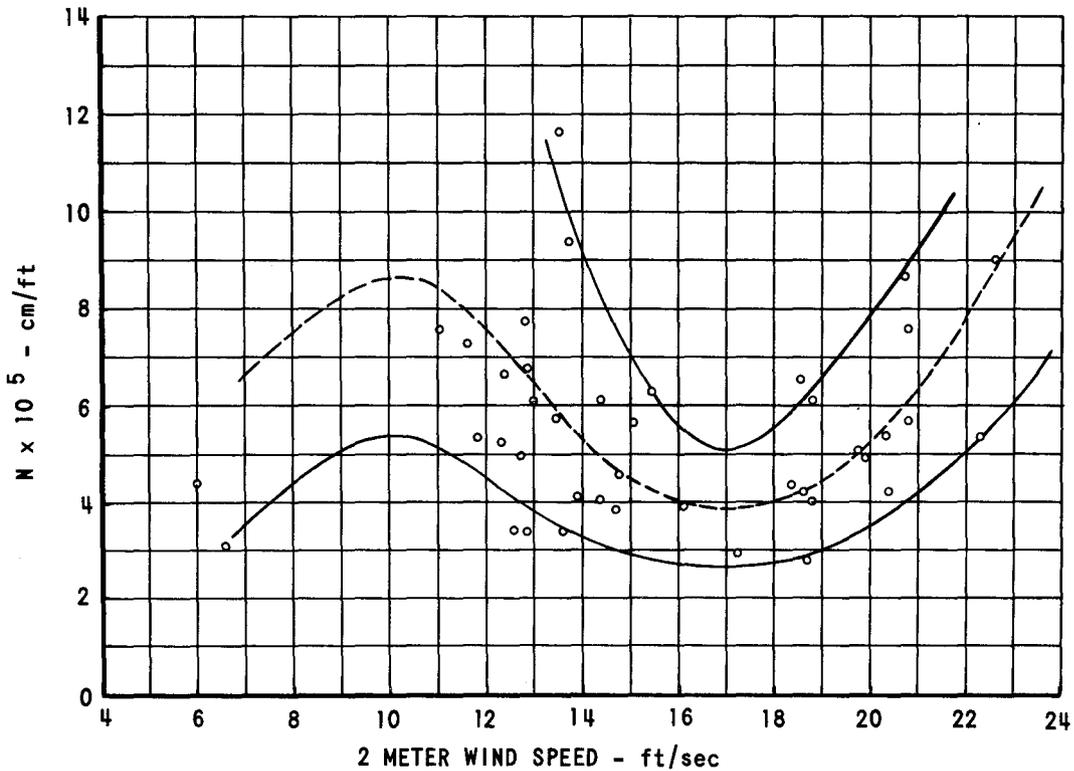
The function $P(V_2)n(V_2)$ was plotted as a function of V_2 and the integral evaluated using a planimeter. With the value of the integral computed, K was then determined from equation (37). The curve

$$K + n(V_2) = N(V_2)$$

(figure 32) is the adjusted curve and represents, as closely as we are able to determine, the variation of N with wind speed (and waves). The solid line of figure 32 is the final form of the $N(V)$ relationship determined by the CAL study. The dotted line of figure 32 is the water budget mean value of N .



a) Mid-Lake Tower - 2 Meter Data.



b) Intake Tower - 2 Meter Data

Figure 31.—Mass transfer coefficient N as a function of wind speed.

The $N(V)$ curve presented in figure 32 is the "instantaneous" value of N for the equilibrium wind speed V . The term "instantaneous" refers to the length of time the wind speed, and other parameters required in the evaporation estimate, remain rela-

tively constant. The curve was developed using mean measurements over 10-minute time intervals.

It should be emphasized that the variable N relationship shown in figure 32 should *not* be used with long-period averaged inputs. For example, if the wind

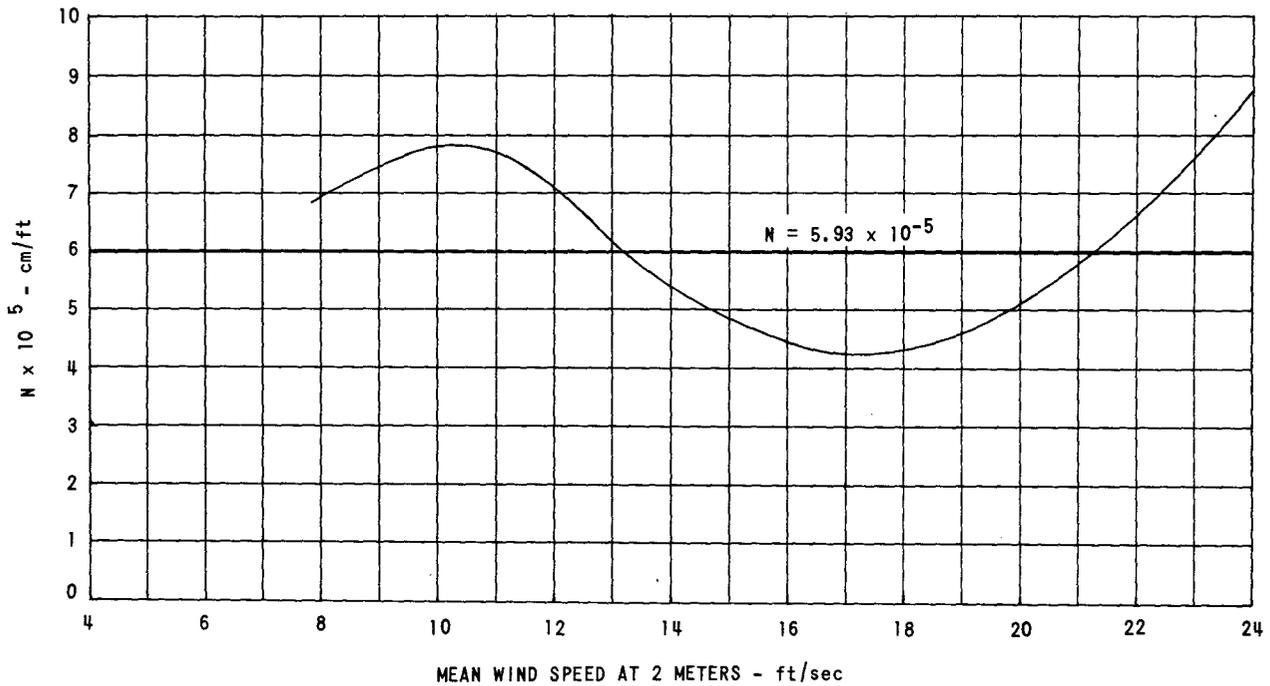


Figure 32.—Mass transfer coefficient N as a function of wind speed—curve adjusted to fit water budget data.

speed were constant at 13 feet per second for 12 hours and then increased to 21 feet per second for 12 hours, the mean wind for the 24-hour period would be 17 feet per second. If we now go to the curve with this 24-hour mean wind, we find that N is 4.3×10^{-5} , a low and incorrect value. If, on the other hand, we recognize the true meaning of the curve, we must determine N for 13 feet per second and again for 21 feet per second. In both cases this value is about 6.0×10^{-5} , making the correct value for the 24-hour period 6.0×10^{-5} .

If the curve is correctly applied, it should lead to more accurate estimates of evaporation over long time periods where wind speeds have remained relatively constant during certain portions of the period. For example, if during a one-week survey period, the

wind speed remained near 17 feet per second for a large percentage of the time, use of the water-budget-derived coefficient would lead to an overestimate of the actual evaporation. Figure 32 shows that the true evaporation would be approximately 30 percent less than that given by the constant- N relationship. To successfully make use of this new relationship for long survey periods it will generally be necessary to compute the evaporation for short periods (possibly one hour) and sum the results. For very long periods, however, where one can expect the wind speed distribution to approach the one used in the analysis, the new $N(V)$ curve will not improve the estimate over the $N = \text{constant}$ approach. In fact, the two methods should give the same result since the $N(V)$ curve was adjusted to satisfy just this requirement.

CONCLUSIONS AND RECOMMENDATIONS

Despite certain difficulties with techniques and instrumentation, some very useful results have been obtained. The wave tank measurements showed that waves apparently modify the air flow over the surface in such a way that turbulent transport mechanisms can be suppressed. Follow-up experiments conducted in the field tend to support the tank measurements, and in fact show that the evaporation retarding mechanism appears to be even more effective in the field.

The results of both the laboratory and the field studies tend to confirm the postulates of Stewart and the theory of Phillips outlined in the Introduction. Organized flow does develop in the wave troughs, as demonstrated by the flow tracing experiments in the wave tank. The consistent minima in the mass transport coefficient curves, developed from the Lake Hefner data, tend to show a resonance effect for a certain combination of wind speed and H/T ratio.

No information was gained about the effect of breaking bubbles on evaporation rate. However, in view of the measurements made, it would seem that the controlling mechanism for evaporation from a free water surface is the turbulent transport mechanism, at least at light-to-moderate wind speeds. In a wind regime where bubbling on the waves might be important, the wind-wave mechanism is in control and vertical transport of moisture is being retarded. At high wind speeds, spray and bubbles may well be important.

Although the work performed by CAL was not designed to evaluate evaporation-suppressing chemicals, use of these chemicals during certain periods of the Lake Hefner study was evident in the processed data. Curiously enough, the mass transport coefficients computed when the alcohol film was near or around the measurement site, while being somewhat lower than the main bulk of data, displayed a similar dependence on wind and waves. We might conclude from this that the existence of the monolayer does not significantly alter the wave parameters that influence the transport mechanisms. Unfortunately, we were not able to obtain a significant sample of wave spectra during periods of film application to make any judgment on the effect of monolayers on waves.

A mass transfer coefficient curve was developed which is applicable to short-period measurements and takes into account the wind-wave interaction. This type of relationship will be valuable in future work related to the evaluation of evaporation-retarding chemicals where it will be most desirable to have a method of estimating short-term evaporation rates from the untreated surface.

Recommendations

While the results of the present study are encouraging, much is yet to be learned about the air-wave interaction. The results obtained at Lake Hefner are not necessarily applicable to other lakes where different wind-wave characteristics may be present. This is especially true of larger lakes where the fetch is much longer than on Lake Hefner.

It is therefore recommended that air-wave interaction studies be continued with specific areas of attention as follows:

1. Develop wind-wave and wind-evaporation rate studies on a lake much larger than Hefner in order to better assess the effect of fetch.
2. Attempt to determine the actual air flow which develops over waves and formulate an analytical description of this flow.
3. Determine the effect of monolayers on the wave regime. There may be some interference effect such as shifting the wave parameters out of the natural resonance point.
4. Increase the accuracy of the eddy flux evaporation measurement technique and thus reduce the scatter in the computed data. With increased accuracy, the mass transfer characteristic could be computed directly without the necessity of adjustment to long-period measurements which may have unknown inherent biases.
5. Extend the wind-wave and wind-evaporation studies to much higher wind speeds in order to deduce: (a) the effect of spray on evaporation rates, and (b) the existence of higher "harmonics" of the null effect found in the mass transport coefficient at 17 feet per second wind speed.

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APPENDIX A

Bridge Circuit Used for Thermistor Temperature Measurements in Wind Tunnel

The bridge is driven by a low amplitude, repetitive, 1-millisecond pulse from the pulse generator. When the bridge is unbalanced, a signal is developed across the primary winding of T_1 and fed to the amplifier. The amplifier output is fed to a balanced detector circuit through a split-winding transformer. If the bridge is unbalanced in one direction, a positive voltage is developed across C_1 , unbalancing the meter. If the bridge is unbalanced in the other

direction, the positive voltage is developed across C_2 and the meter reverses. A zero-center meter is used. Sensitivity is determined by the amount of gain set in the amplifier and can be very high. The accuracy of balance in the equipment used is $\pm 0.1^\circ \text{C}$. Balancing is done with a 10-turn potentiometer, (R_3), equipped with a digital readout dial. Balancing of the bridges takes only a couple of seconds and the digital readout makes the recording task very simple.

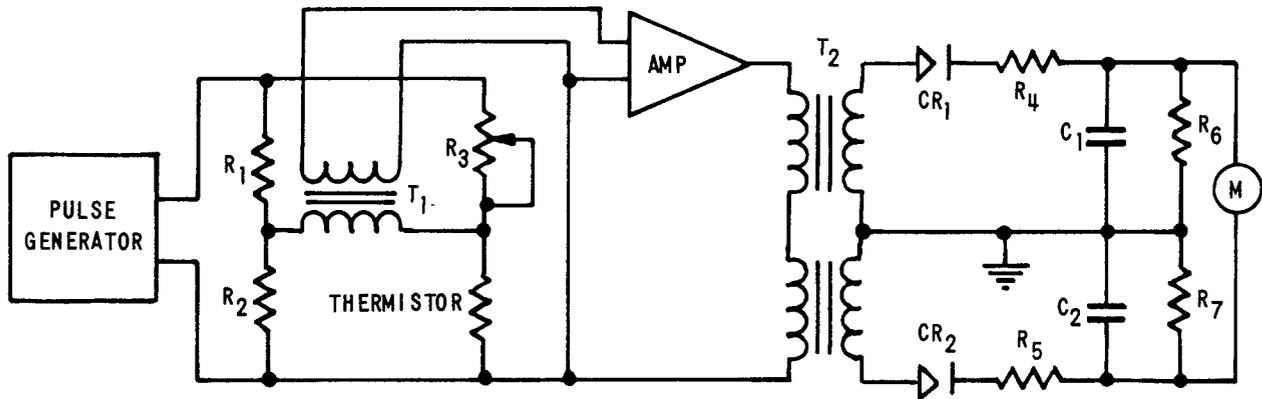


Figure A-1.—Schematic diagram of pulsed thermistor bridge.

APPENDIX B

Solution of Wave Tank Electrical Analog

The electrical analog for the wave tank experiment is given below.

The problem is to find the time function of the charge on C_2 after closing switch S . The theoretical amplifier A in the diagram has infinite input impedance, zero output impedance, and unity gain. The perfect isolation of the amplifier then permits us to solve the above circuit in two stages. First we solve for the voltage on C_1 , then use this relationship as the driving voltage for the second circuit involving R_3 and C_2 .

Upon closing switch S we may write for the voltage around the first loop

$$(i_1 + i_2)R_1 + \frac{q}{C_1} - V = 0 \quad (\text{B-1})$$

where $i_1 = q/C_1R_2$

Then equation (B-1) becomes

$$\left(\frac{q}{C_1R_2} + \frac{dq}{dt} \right) R_1 + \frac{q}{C_1} - V = 0.$$

Differentiating with respect to t

$$\frac{R_1}{C_1R_2} \frac{dq}{dt} + R_1 \frac{d^2q}{dt^2} + \frac{1}{C} \frac{dq}{dt} = 0$$

Rearranging terms and simplifying we have

$$\frac{d^2q}{dt^2} + \frac{R_1 + R_2}{R_1R_2C_1} \frac{dq}{dt} = 0 \quad (\text{B-2})$$

Let $\frac{R_1}{R_1 + R_2} = r$ and $R_1C_1 = \tau$

Then equation (B-2) becomes

$$\frac{d^2q}{dt^2} + \frac{1}{r\tau} \frac{dq}{dt} = 0 \quad (\text{B-3})$$

This equation has a solution of the form

$$q(t) = Ae^{-(t/\tau r)} + B$$

Let us set the boundary conditions

$$t=0, \quad q=q_0 \quad \text{and} \quad t=\infty, \quad q=C_1Vr$$

Thus $A = q_0 - B$ and $B = CVr$.

The final solution is then

$$\begin{aligned} q(t) &= (q_0 - CVr)e^{-(t/\tau r)} + CVr \\ &= q_0e^{-(t/\tau r)} + CVr(1 - e^{-(t/\tau r)}) \end{aligned}$$

Dividing by C_1 we have

$$V(t) = V_0e^{-(t/\tau r)} + Vr(1 - e^{-(t/\tau r)}) \quad (\text{B-4})$$

We may now use equation (B-4) for the voltage

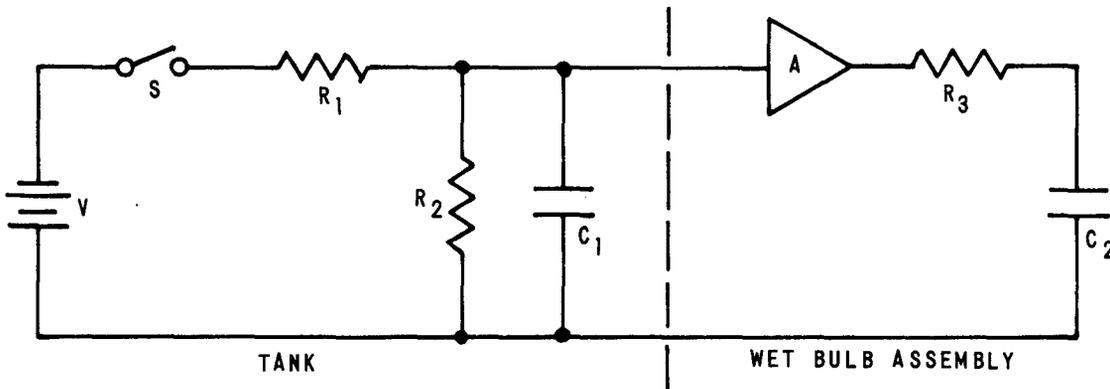


Figure B-1.—Electrical analog of evaporation experiment.

applied to R_3C_2 . Equating the voltages around the second loop we have

$$R_3 \frac{dq}{dt} + \frac{q}{C_2} = \frac{q_0}{C_2} e^{-(t/r\tau)} + \frac{q_m}{C_2} (1 - e^{-(t/r\tau)})$$

where $V_0 = \frac{q_0}{C_2}$ and $Vr = \frac{q_m}{C_2}$.

Dividing by R_3 we have

$$\frac{dq}{dt} + \frac{q}{\tau_m} = \frac{q_0}{\tau_m} e^{-(t/r\tau)} + \frac{q_m}{\tau_m} (1 - e^{-(t/r\tau)}) \quad (\text{B-5})$$

where $\tau_m = R_3C_2$.

Equation (B-5) may be integrated by using the integrating factor $e^{(t/\tau_m)}$. Multiplying both sides of equation (B-5) by $e^{(t/\tau_m)}$ we have

$$\left(\frac{dq}{dt} + \frac{q}{\tau_m} \right) e^{(t/\tau_m)} = \frac{(q_0 - q_m)}{\tau_m} e^{[t(r\tau - \tau_m)/r\tau\tau_m]} + \frac{q_m}{\tau_m} e^{(t/\tau_m)}$$

$$qe^{(t/\tau_m)} = \frac{q_0 - q_m}{\tau_m} \int e^{[t(r\tau - \tau_m)/r\tau\tau_m]} dt$$

$$+ \frac{q_m}{\tau_m} \int e^{(t/\tau_m)} dt + K$$

$$q(t) = \frac{r\tau(q_0 - q_m)}{r\tau - \tau_m} e^{[t(r\tau - \tau_m)/r\tau\tau_m]} e^{-(t/\tau_m)}$$

$$+ q_m + Ke^{-(t/\tau_m)}$$

$$q(t) = q_m + \frac{r\tau(q_0 - q_m)}{r\tau - \tau_m} e^{-(t/r\tau)} + Ke^{-(t/\tau_m)}$$

(B-6)

Apply boundary condition $q(t) = q_0$ when $t = 0$

$$K = q_0 - q_m - \frac{r\tau(q_0 - q_m)}{r\tau - \tau_m} \quad (\text{B-7})$$

Combining equations (B-6) and (B-7) and ar-

ranging terms we have

$$q(t) = q_m + \frac{(q_0 - q_m)r\tau}{r\tau - \tau_m} e^{-(t/r\tau)} + \frac{(q_m - q_0)\tau_m}{r\tau - \tau_m} e^{-(t/\tau_m)} \quad (\text{B-8})$$

Equation (B-8) is the solution of the electrical circuit where q is the charge on the Capacitors C_2 .

Going to the tank experiment, q becomes the specific humidity and the boundary conditions are

q_0 = equilibrium specific humidity at start of experiment

q_s = saturation specific humidity at the water temperature

q_m = maximum specific humidity reached at $t \rightarrow \infty$

$r = q_m/q_s$.

The other terms in equation (B-8) become

τ = actual time constant of experiment if no air leakage were present

τ_m = time constant of wet bulb.

Equation (B-8) thus becomes the solution to the evaporation experiment. Adding $(q_m - q_0)$ to both sides of equation (B-8) and rearranging terms we have

$$q(t) = q_0 + (q_m - q_0) \left[1 - \frac{r\tau}{r\tau - \tau_m} e^{-(t/r\tau)} + \frac{\tau_m}{r\tau - \tau_m} e^{-(t/\tau_m)} \right]$$

Taking q_0 to the left side of the equation and dividing through by $(q_m - q_0)$ we have

$$\frac{q - q_0}{q_m - q_0} = 1 - \frac{r\tau}{r\tau - \tau_m} e^{-(t/r\tau)} + \frac{\tau_m}{r\tau - \tau_m} e^{-(t/\tau_m)}$$

and therefore

$$\left[1 - \frac{q - q_0}{q_m - q_0} \right] = \frac{r\tau}{r\tau - \tau_m} e^{-(t/r\tau)} - \frac{\tau_m}{r\tau - \tau_m} e^{-(t/\tau_m)} \quad (\text{B-9})$$

Equation (B-9) is in a normalized form where only time constants are important. The start time and end points do not displace the experimental curve if the time constants remain fixed. This form permits much better comparison of the experimental results.

APPENDIX C

Design and Construction Details of the Wave Probe

The expression for the capacitance of a cylindrical condenser is given by

$$C = \frac{2.4 \times 10^{-7} \epsilon L}{\log b/a} \mu \text{ farads} \quad (\text{C-1})$$

where ϵ = dielectric constant

L = length of electrodes

a = diameter of inner conductor

b = diameter of outer conductor.

From equation (C-1) we see that the capacitance is a linear function of the length of the conductors. In our apparatus, the water level on the wave probe determines the length of the outer conductor and hence the length of the capacitor. With a 3-inch separation of electrodes, the stray capacity between them above the water line is extremely small and can be neglected for the purpose of design. The capacitance of the probe is then a linear function of the water level on the insulated conductor. It remains to transform this capacitance value into a voltage and still retain the linear relationship to the water level.

A block diagram of the system used to make this transformation is shown in figure C-1. An AC signal is fed to the probe through series resistor R . The impedance of the probe and the resistance R determines the current, i , that flows in the input loop. The voltage, V_P , that is developed across the resistance, R , is then a function of the current flow, i , which in turn

is a function of the water level on the probe. This voltage is given by

$$V_P = iR = \frac{V_i R}{\sqrt{(R+R_w)^2 + (1/C\omega)^2}}$$

$$= \frac{V_i R \omega C}{\sqrt{[(R+R_w)C\omega]^2 + 1}} \quad (\text{C-2})$$

where $\omega = 2\pi f$

f = frequency of AC drive signal

C = capacity of probe

V_i = input drive source voltage.

Equation (C-2) gives the probe output voltage in terms of the other parameters involved. We see that this expression is in general not linear with C . However, if we choose the appropriate values of R and f and make the physical dimensions of the probe such that the maximum value of C is not too large, we can keep the magnitude of the term $[(R+R_w)C\omega]^2$ small (< 1). As long as the maximum value of this term is less than unity, the denominator in equation (C-2) is very nearly constant. Equation (C-2) then may be written

$$V_P \approx KC \quad (\text{C-3})$$

where K is a constant, with the accuracy of the approximation being under the control of the designer. The values actually chosen for the wave

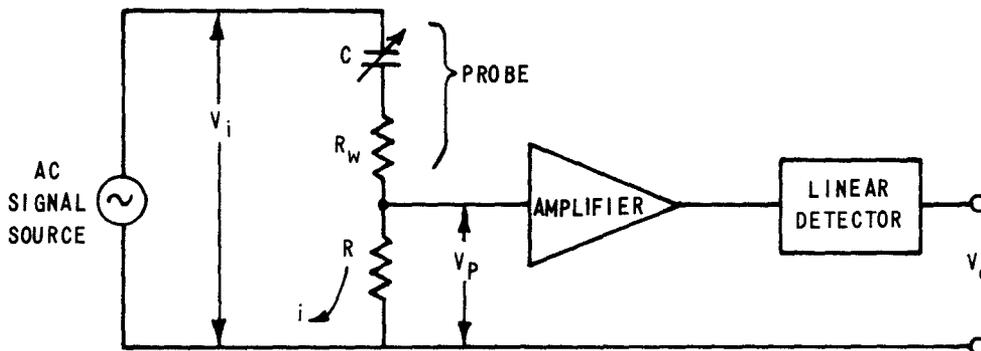


Figure C-1.—Block diagram of probe system.

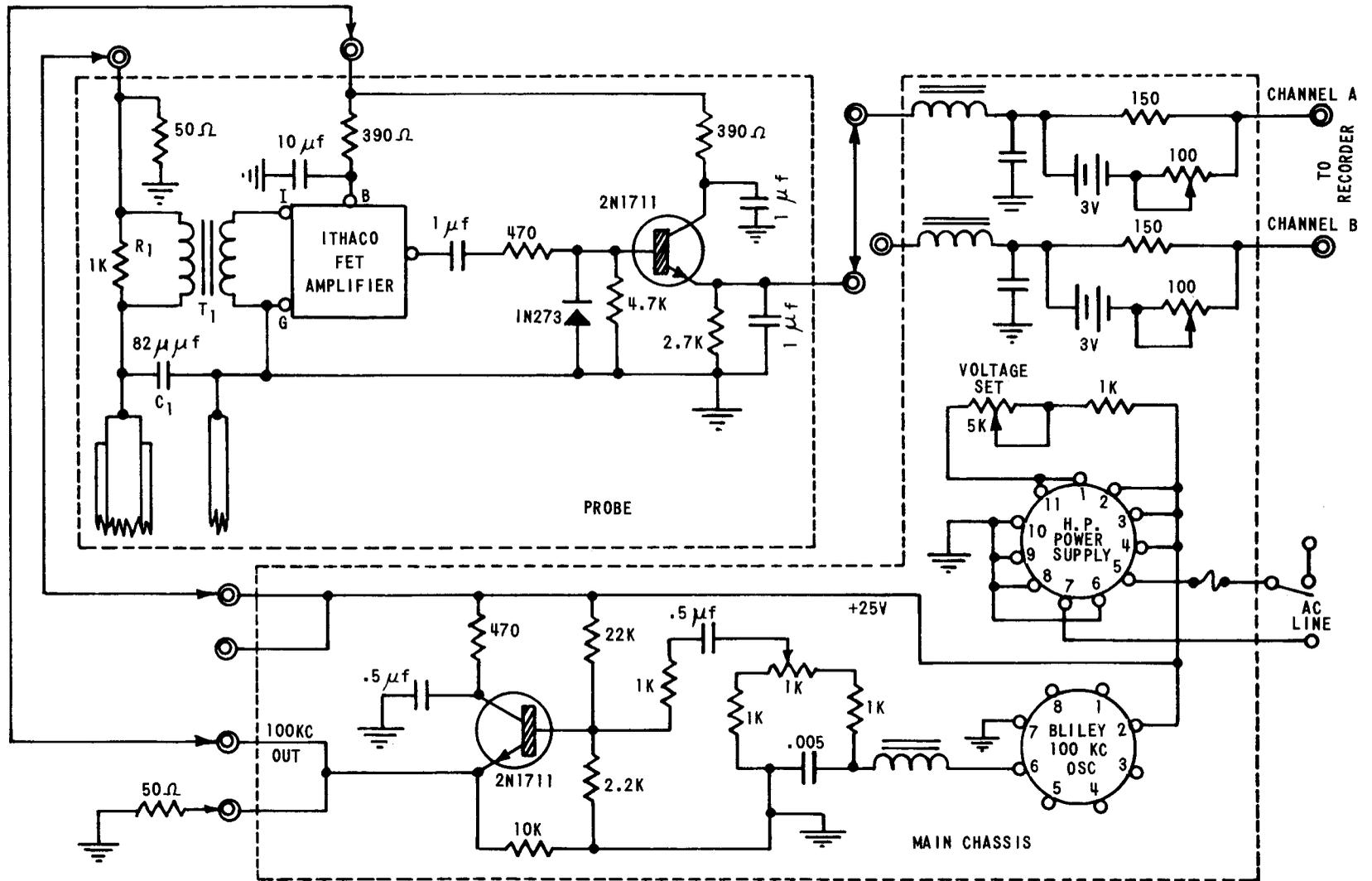


Figure C-2.—Schematic diagram of wave probe system.

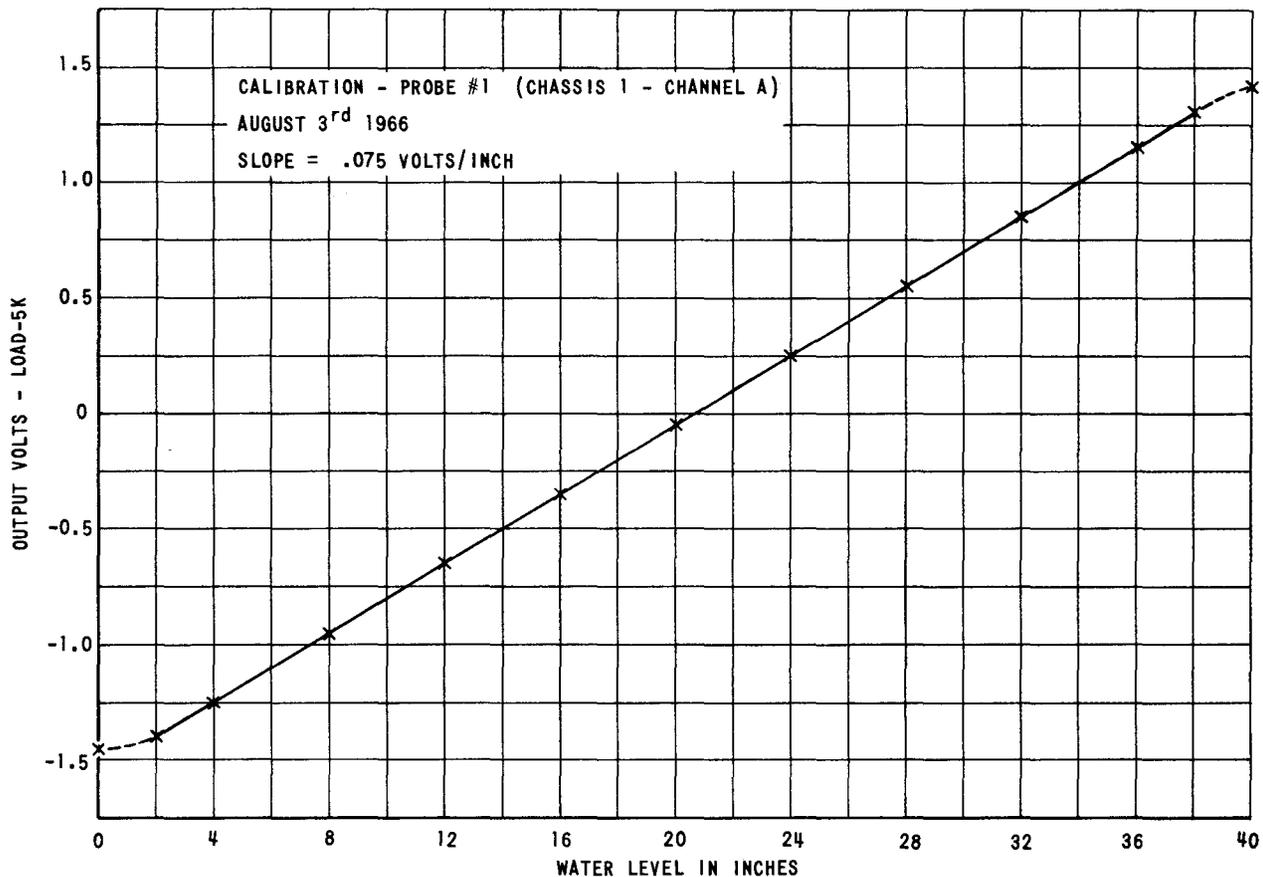


Figure C-3.—Wave probe calibration curve.

probe are as follows:

$$R = 1000 \text{ ohms}$$

$$C_{\text{max}} = 500 \mu\mu \text{ farads (full length of probe)}$$

$$f = 100 \text{ kilohertz.}$$

The resistance of the water (R_w) was found to be negligible in comparison with the reactance of the capacitor. With these values, the denominator in equation (C-2) has a maximum value of 1.05. Since the minimum value is unity, the term is essentially a constant and equation (C-3) is valid.

The probe voltage V_P is fed into an amplifier with a high input impedance (approximately 300 kilohms) so that the probe loop is unaffected by the external circuits. The 100-kilohertz signal is amplified and detected in a linear detector circuit. The input voltage V_i and the amplifier gain were chosen so that an output voltage range of 0–3 volts is obtained over the full immersion range of the probe (about 4 feet). A schematic diagram of the complete probe elec-

tronics is given in figure C-2. Note that in the actual circuit, transformer coupling is used between the probe current sensing resistor, R , and the amplifier. This was done in order to maintain the water at ground potential and minimize noise pick-up. The battery bias source in the output circuit permits centering of output voltage range and is used only to make maximum use of the magnetic tape recorder's dynamic range. The small shunt capacitance, C_1 , across the probe electrodes tunes out the probe inductance at 100 kilohertz.

Figure C-4(a) shows a complete probe with its associated electronics. The white electrode on the left is the insulated probe which, with the water, forms the capacitor. It consists of a one-half-inch brass rod which has been inserted into a five-eighths-inch outer diameter by one-half-inch inner diameter teflon tube. Teflon was chosen for the dielectric material because of its non-wetting properties (less hysteresis effect from water clinging to the surface of probe). The second electrode is an exposed one-half-

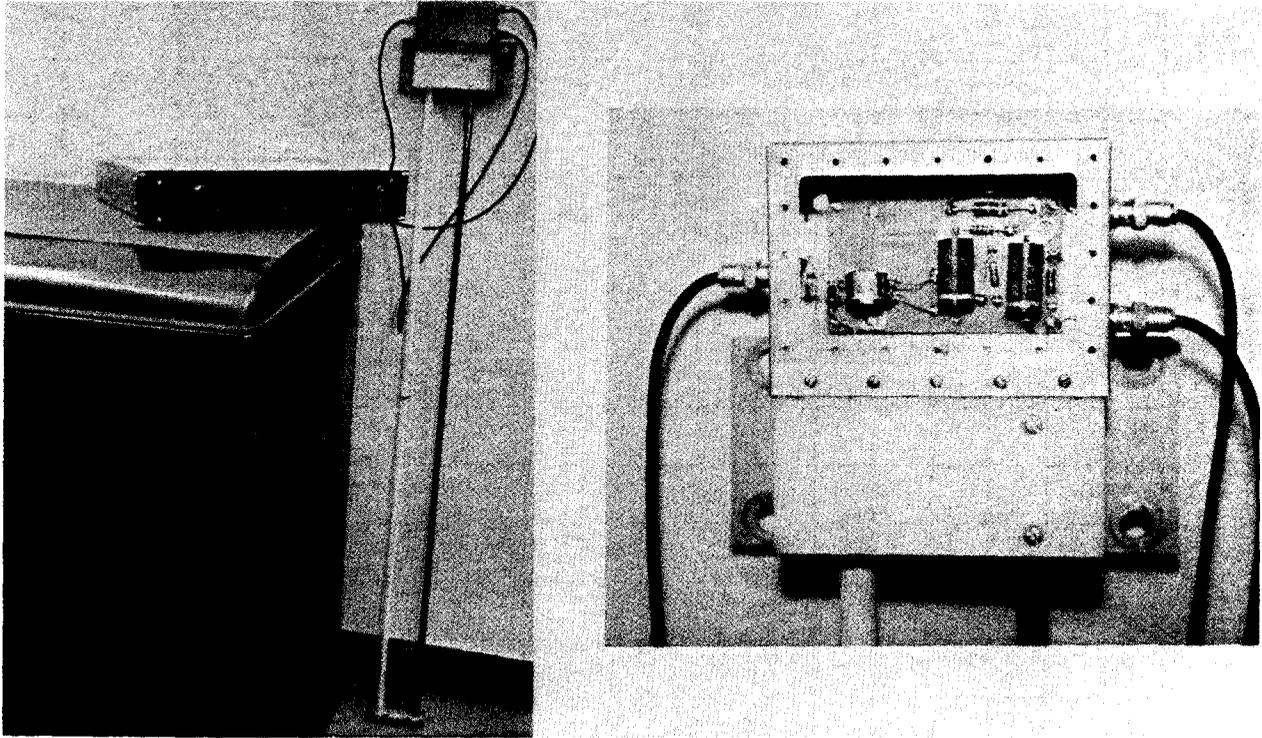


Figure C-4.—Wave measuring probe.

inch brass rod. Figure C-4(b) shows a close-up of the probe head. The two rods are threaded through a 3- by 3-inch solid nylon block for strength and low capacitive strays. A brass box is sealed over the top of the nylon block and houses some of the associated

probe electronics. A rubber gasket around the face plate of the box makes the whole probe head completely water tight. In the field installation, the *BNC* connectors were “potted” in place with a silicone rubber compound.

APPENDIX D

Basic Data From the Wave Tank Evaporation Experiments

V = average wind speed in the evaporation section of the tank in feet per second
 H = average wave height (peak to trough) in inches
 L = wave length in inches
 T = wave period in seconds
 H/L = wave height to length ratio (dimensionless)
 H/T = wave height to period ratio in inches per second

τ = time constant in seconds of experiment, computed from the semilog plots of the data points

D = evaporation coefficient = $v/A\tau$ feet per second where the tank parameter $v/A = 5.8$ feet.

NOTE: Where wave height H is indicated to be zero, small wind-driven capillaries were present. However, no paddle generated waves were present.

Run No.	V	H	L	T	H/L	H/T	τ	D	D/V
32	8.34	0					88	0.066	0.0079
33	6.20	0					113	0.052	0.0084
34	4.16	0					115	0.036	0.00865
35	2.92	0					218	0.027	0.0093
36	10.0	6	120	1.84	0.050	3.26	80	0.073	0.0073
37	6.0	5	120	1.84	0.040	2.72	115	0.051	0.0085
38	4.16	5	120	1.84	0.040	2.72	133	0.044	0.0106
39	2.92	5	120	1.88	0.040	2.72	160	0.036	0.0123
40	12.5	6	80	1.40	0.075	4.30	73	0.080	0.0064
41	8.3	6	80	1.40	0.075	4.30	101	0.058	0.0070
42	5.5	6	80	1.40	0.075	4.30	118	0.049	0.0089
43	3.8	6	80	1.40	0.075	4.30	133	0.044	0.0116
44	15.8	7.5	75	1.20	0.10	6.25	65	0.091	0.0058
45	10.0	8	60	1.44	0.133	7.02	73	0.081	0.0081
46	5.0	8	60	1.1	0.133	7.28	82	0.072	0.0146
47	2.92	8	65	1.19	0.123	6.72	102	0.058	0.0200
48	9.0	3	160	2.35	0.019	1.27	80	0.074	0.0082
49	5.84	2.5	160	2.35	0.016	1.07	109	0.054	0.0096
50	3.0	2	160	2.35	0.013	0.085	162	0.037	0.0123
51	13.0	6	80	1.39	0.075	4.3	74	0.086	0.0066
52	9.0	6	80	1.39	0.075	4.3	97	0.064	0.0071
53	3.0	6	80	1.39	0.075	4.3	149	0.042	0.0140
54	3.0	6	80	1.39	0.075	4.3	158	0.037	0.0123
55	13.0	4	55	0.97	0.073	4.1	69	0.084	0.0065
56	10.8	4	55	0.97	0.073	4.1	84	0.069	0.0064
57	4.0	5	55	0.97	0.091	5.15	118	0.049	0.0122
58	8.3	3	145	2.1	0.021	1.43	84	0.069	0.0083
59	6.7	2	145	2.1	0.014	0.95	104	0.056	0.0084
60	2.5	2	145	2.1	0.014	0.95	188	0.031	0.0124
61	10.0	4	80	1.39	0.050	2.88	74	0.079	0.0079
62	8.3	4	80	1.39	0.050	2.88	88	0.066	0.0079
63	3.0	4	80	1.39	0.050	2.88	152	0.038	0.0126
64	13.3	6.5	60	1.06	0.110	6.13	75	0.078	0.0059
65	9.2	6	60	1.06	0.10	5.66	97	0.060	0.0065
66	3.5	6	60	1.06	0.10	5.66	136	0.043	0.0123

Run No.	V	H	L	T	H/L	H/T	τ	D	D/V
67	8.3	2	145	2.3	0.014	0.87	85	0.069	0.0083
68	6.0	2	145	2.3	0.014	0.87	112	0.052	0.0087
69	2.5	3.5	145	2.3	0.024	1.57	204	0.029	0.0116
70	10.0	3.5	80	1.33	0.044	2.64	78	0.075	0.0075
71	5.0	0					114	0.051	0.0102
72	7.5	4	80	1.33	0.050	3.0	103	0.057	0.0076
73	2.5	4	80	1.33	0.050	3.0	180	0.032	0.0128
74	13.0	6	55	0.90	0.11	6.68	80	0.073	0.0056
75	8.3	6	55	0.90	0.11	6.68	98	0.059	0.0071
76	3.0	6	55	0.90	0.11	6.68	154	0.038	0.0126
77	14.0	4	50	0.83	0.08	4.83	69	0.084	0.0060
78	10.0	5	50	0.83	0.10	6.02	84	0.069	0.0069
79	3.5	4.5	50	0.83	0.090	5.4	118	0.049	0.0140
80	12.5	0					65	0.089	0.0071
81	13.3	4	150	2.36	0.027	1.69	59	0.098	0.0074
82	14.0	4.5	120	1.70	0.038	2.65	56	0.104	0.0074
83	17.0	6	85	1.39	0.071	4.30	53	0.110	0.0065
84	20.0	6.5	65	1.2	0.10	5.4	47	0.123	0.00615
85	20.0	6.5	65	1.2	0.10	5.4	46	0.126	0.0063
86	12.5	4	150	2.41	0.027	1.66	56	0.104	0.0083
87	13.3	4	110	1.70	0.036	2.35	58	0.10	0.0075
88	15.0	5.5	80	1.39	0.069	3.96	54	0.107	0.0071
89	20.8	6	60	1.0	0.1	6.0	48	0.121	0.0058
90	11.7	2	150	2.20	0.013	0.91	60	0.097	0.0083
91	12.5	3.5	120	1.61	0.029	2.18	58	0.100	0.0080
92	15.8	5	80	1.33	0.063	3.76	51	0.114	0.0072
93	20.8	6.5	60	1.0	0.11	6.5	51	0.114	0.0055
94	12.0	2	150	2.2	0.013	0.91	56	0.104	0.0087
95	12.5	2.5	120	1.70	0.021	1.47	62	0.094	0.0075
96	16.0	4.5	75	1.50	0.060	3.0	51	0.114	0.0071
97	18.4	6	60	1.11	0.10	5.4	50	0.116	0.0063
98	15.8	0					47	0.123	0.0078
99	15.8	3	150	2.12	0.020	1.42	50	0.116	0.0074
100	16.7	4	110	1.6	0.036	2.5	46	0.126	0.0076
101	18.0	4	75	1.3	0.053	3.06	43	0.135	0.0075
102	25.0	5.5	50	0.87	0.11	6.32	38	0.152	0.0061
103	16.0	4	150	2.4	0.027	1.67	45	0.129	0.0081
104	18.0	4	115	1.68	0.035	2.38	40	0.145	0.0081
105	20.0	5.5	80	1.39	0.069	3.96	40	0.145	0.0073
106	15.8	5.5	155	2.35	0.036	2.34	46	0.126	0.0080
107	18.5	6	120	1.93	0.050	3.11	40	0.145	0.0078
108	21.0	6	80	1.39	0.075	4.33	41	0.141	0.0067

APPENDIX E

Reduced Lake Hefner Data

Key to column headings

- V = wind speed in meters per second
- E = vertical flux of water vapor in grams per meter squared per second
- $\Delta q = (q_s - q_z)$ = difference in specific humidity between the measured value at height z and the saturation value at the water surface temperature—grams per gram

$$D = E / \rho \Delta q \text{— feet per second}$$

$$n = \frac{\rho D}{V} \text{ grams per centimeter squared per}$$

foot $\times 10^{-5}$ (mass transfer coefficient)

σ^2_{WH} = variance of wave height record in inches²

H/T = wave parameter, height to period ratio—feet per second.

Key to time of Sample No.

Time shown is the beginning time of sample.

Sample No.	Date	Time C.S.T.	Sample No.	Date	Time C.S.T.
1	8/25/66	2032	32	8/30/66	0158
2	"	2358	33	"	0458
3	8/26/66	0628	34	"	0958
4	"	1258	35	"	1258
5	"	1758	36	"	1458
6	"	2358	37	"	1758
7	8/27/66	0753	38	"	2238
8	"	1158	39	8/25/66	1815
9	"	1558	40	"	2000
10	8/28/66	1958	41	"	2016
11	"	0158	42	"	2034
12	"	0458	43	8/26/66	1050
13	"	0828	44	"	1200
14	"	1158	45	"	1400
15	"	1628	46	"	1800
16	"	1728	47	"	1822
17	"	1828	48	"	1924
18	"	1958	49	"	1956
19	"	2328	50	"	1902
20	8/29/66	0058	51	8/27/66	1112
21	"	0447	52	8/25/67	1621
22	"	0613	53	"	1645
23	"	0913	54	"	1700
24	"	0948	55	"	1716
25	"	1348	56	"	1732
26	"	1443	57	8/29/67	1521
27	"	1559	58	"	1537
28	"	1649	59	"	1552
29	"	1752	60	"	1608
30	"	2058	61	"	1623
31	"	2328	62	"	1636

MID-LAKE TOWER—2 meters

Sample No.	V	E	Δq	D	n	σ^2_{WH}	H/T
1							
2	3.71	0.0236	0.0064	0.0105	3.55	1.10	0.23
3	3.88	0.0219	0.0099	0.0062	2.01	1.25	0.23
4	6.33	0.0273	0.0089	0.0089	1.77	3.0	0.31
5	5.79	0.0217	0.0101	0.0062	1.35	2.6	0.29
6	6.35	0.0301	0.0016	0.0053	1.05	2.9	0.31
7	6.10						
8	5.50	0.0272	0.0074	0.0105	2.17	2.35	0.28
9	5.44	0.0151	0.0091	0.0048	1.10	2.30	0.28
10	4.68	0.0245	0.0098	0.0072	1.94	1.80	0.26
11	5.79	0.0362	0.0068	0.0151	3.27	2.60	0.29
12	5.20	0.0179	0.0056	0.0091	2.20	2.20	0.27
13	6.81	0.0341	0.0051	0.0192	3.56	2.90	0.288
14	3.86	0.0180	0.0053	0.0098	3.20	1.25	0.23
15	4.40	0.0035	0.0063	0.0016	0.46	1.60	0.25
16	6.44	0.0018	0.0061	0.00085			
17	3.87	0.0145	0.0044	0.0095	3.10	1.25	0.23
18	4.11	0.0045	0.0053	0.0025	0.49	1.40	0.24
19	5.43	0.0106	0.0056	0.0054	1.25	2.35	0.28
20	5.54	0.0083	0.0065	0.0037	0.84	2.40	0.28
21	4.69	0.0130	0.0062	0.0060	1.62	1.80	0.26
22	5.34	0.0136	0.0064	0.0061	1.44	2.30	0.28
23	5.28	0.0082	0.0054	0.0044	1.05	2.20	0.27
24	4.89	0.0110	0.0058	0.0054	1.38	1.90	0.26
25	3.57	0.0080	0.0064	0.0036	1.26	1.10	0.22
26	7.61	0.0270	0.0072	0.0108	1.78	4.0	0.34
27	2.97	0.0034	0.0066	0.0015	0.63	0.70	0.21
28	5.56	0.0075	0.0078	0.0028	0.63	2.40	0.28
29	4.91	0.0066	0.0051	0.0037	0.95	1.95	0.26
30	3.68	0.0047	0.0070	0.0019	0.635	1.15	0.23
31	3.70	0.0067	0.0053	0.0036	1.22	1.20	0.23
32	3.59	0.0061	0.0042	0.0042	1.43	1.10	0.22
33	2.17	0.0038	0.0045	0.0024	1.39	0.40	
34	3.11	0.0079	0.0062	0.0038	1.53	0.75	0.21
35	4.64	0.0160	0.0053	0.0088	2.38	1.80	0.25
36	5.13	0.0074	0.0055	0.0039	0.96	0.23	
37	5.89	0.0201	0.0054	0.0108	2.30	2.65	0.29
38	4.20	0.0108	0.0050	0.0063	1.88	1.40	0.24
39							
40	2.77	0.0769	0.0076	0.0290	13.2	0.45	0.198
41	2.74	0.0186	0.0073	0.0073	3.34	0.45	0.205
42							
43	7.34	0.0338	0.0107	0.0088	1.53	3.91	0.33
44	6.51	0.0140	0.0111	0.0036	0.71	2.74	0.29
45	7.02	0.0386	0.0100	0.0110	2.0	5.70	0.32
46	5.22	0.0323	0.0111	0.0084	2.0	2.34	0.27
47	4.31	0.0373	0.0107	0.0100	2.94	1.57	0.24
48	2.88	0.0276	0.0105	0.0075	3.24	0.75	0.20

MID-LAKE TOWER—2 meters—Continued

Sample No.	V	E	Δq	D	n	σ^2_{WH}	H/T
49							
50	3.68	0.0141	0.0105	0.0039	1.32	1.07	0.23
51	6.09	0.0180	0.0078	0.0066	1.36	3.08	0.32
52	4.07	0.0177	0.0142	0.0037	1.12	1.35	0.24
53	4.24	0.0490	0.0146	0.0098	2.88	1.50	0.24
54	4.50	0.0463	0.0146	0.0093	2.60	1.70	0.25
55	3.85	0.0586	0.0141	0.0122	4.0	1.20	0.23
56	4.11	0.0426	0.0144	0.0086	2.64	1.35	0.24
57	2.29	0.0639	0.0104	0.0181	9.9	0.40	0.18
58	2.72	0.0833	0.0105	0.0234	10.0	0.60	0.20
59	2.32	0.107	0.0112	0.0282	15.2	0.50	0.19
60	2.70	0.0913	0.0104	0.0260	14.8	0.60	0.19
61	2.18	0.0905	0.0110	0.0243	14.0	0.35	0.18
62	2.16	0.0662	0.0107	0.0183	10.0	0.35	0.18

MID-LAKE TOWER—8 meters

Sample No.	V	E	Δq	D	Sample No.	V	E	Δq	D
1					32	4.22	0.0093	0.0045	0.0059
2	4.16	0.0246	0.0097	0.0072	33	2.72	0.0110	0.0046	0.0068
3	4.34	0.0289	0.0098	0.0083	35	5.73	0.0119	0.0059	0.0058
4	7.74	0.0810	0.0109	0.0215	36	6.28	0.0157	0.0061	0.0075
5	6.86	0.0500	0.0113	0.0128	37	7.04	0.0166	0.0057	0.0085
6	7.48	0.0852	0.0094	0.0258	38	5.01	0.0166	0.0074	0.0065
7	7.16	0.0682	0.0083	0.0234	39				
8	6.53	0.0636	0.0088	0.0207	40	2.74	0.0235	0.0125	0.0053
9	6.22	0.0236	0.0104	0.0066	41	2.84	0.0112	0.0125	0.0025
10	5.30	0.0173	0.0108	0.0046	42				
11	6.71	0.0360	0.0070	0.0146	43	9.15	0.0664	0.0107	0.0177
12	5.74	0.0260	0.0052	0.0142	44	7.85	0.0718	0.0109	0.0187
13	8.03	0.0240	0.0050	0.0138	45	8.80	0.0590	0.0112	0.0149
14	4.54	0.0230	0.0057	0.0117	46	6.26	0.0491	0.0120	0.0118
15	5.34	0.0263	0.0072	0.0106	47	5.34	0.0452	0.0118	0.0110
16	7.52	0.0257	0.0067	0.0111	48	3.46	0.0340	0.0113	0.0086
17	4.57	0.0105	0.0066	0.0046	49				
18	4.72	0.0068	0.0057	0.0034	50	4.39	0.0269	0.0117	0.0066
19	6.30	0.0164	0.0059	0.0079	51	7.18	0.0422	0.0092	0.0131
20	6.25	0.0123	0.0066	0.0053	52				
21	5.30	0.0200	0.0060	0.0095	53				
22	5.86	0.0156	0.0061	0.0073	54				
23	6.15	0.0205	0.0057	0.0102	55				
24	6.04	0.0119	0.0057	0.0060	56				
25	4.58	0.0136	0.0063	0.0062	57	2.57	0.0495	0.0108	0.0137
26	9.23	0.0230	0.0071	0.0092	58	2.96	0.0650	0.0113	0.0172
27	3.86	-0.0022	-0.0061	-0.0010	59	2.61	0.1036	0.0119	0.0260
28	6.70	0.0255	0.0075	0.0133	60	3.46	0.1197	0.0117	0.0305
29	5.58	0.0084	0.0055	0.0044	61	2.56	0.1140	0.0122	0.0280
30	4.47	0.0096	0.0073	0.0037	62	2.56	0.0644	0.0119	0.0162
31	4.39	0.0087	0.0052	0.0048					

INTAKE TOWER—2 meters

Sample No.	V	E	Δq	D	n	σ^2_{WH}	H/T
1	3.37	0.0599	0.0084	0.0203	7.55	1.05	0.19
2	3.98	0.0534	0.0080	0.0190	6.04	1.75	0.23
3	4.11	0.0532	0.0081	0.0186	5.70	2.0	0.24
4	6.30	0.1434	0.0094	0.0435	8.66	5.54	0.33
5	6.05	0.0685	0.0081	0.0243	5.05	5.28	0.32
6	6.89	0.0795	0.0046	0.0493	9.0	6.0	0.34
7	6.16	0.0624	0.0068	0.0202	5.3	5.41	0.33
8	5.63	0.0460	0.0067	0.0185	4.35	4.75	0.31
9	4.68	0.0365	0.0045	0.0233	6.25	3.03	0.27
10	3.95	0.0469	0.0063	0.0212	6.76	1.75	0.23
11	6.34	0.0628	0.0047	0.0381	7.55	5.60	0.33
12	4.97	0.0603	0.0038	0.0452	11.4	3.65	0.29
13	6.33	0.0298	0.0030	0.0283	5.6	5.60	0.33
14	3.91	0.0310	0.0037	0.0241	7.73	1.70	0.23
15	4.38	0.0010	0.0030	0.0009	0.28	2.50	0.26
16	5.57	-0.0031	0.0038	-0.0024			
17	3.74	0.0098	0.0028	0.0101	3.37	1.45	0.21
18	4.41	0.0179	0.0024	0.0215	6.10	2.60	0.26
19	5.75	0.0244	0.0038	0.0183	4.0	4.90	0.32
20	5.66	0.0214	0.0032	0.0191	4.25	4.80	0.32
21	4.49	0.0150	0.0031	0.0138	3.86	2.70	0.26
22	4.41	0.0188	0.0038	0.0141	4.03	2.60	0.26
23	5.24	0.0109	0.0025	0.0124	2.95	4.20	0.30
24	4.91	0.0140	0.0026	0.0154	3.9	3.50	0.28
25	3.94	0.0132	0.0036	0.0105	3.34	1.70	0.23
26	7.42	0.0232	0.0061	0.0109	1.86	6.27	0.34
27	1.83	0.0056	0.0026	0.0063	4.30	0.35	
28	4.61	0.0256	0.0035	0.0212	5.68	2.85	0.27
29	4.03	0.0054				1.90	0.23
30	3.53	0.0330	0.0046	0.0204	7.28	1.25	0.20
31	4.18					2.20	0.24
32	4.16	0.0139	0.0036	0.0011	3.30	2.10	0.24
33	2.00	0.0051	0.0030	0.0049	3.06	0.40	
34	3.77	0.0165	0.0030	0.0158	5.28	1.20	0.22
35	3.56	-0.0006	0.0030	-0.0006		1.06	0.20
36	4.61	0.0021	0.0046	0.0013	0.54	2.90	0.26
37	5.71	0.0147	0.0034	0.0124	2.74	4.90	0.32
38	4.13	0.0493	0.0037	0.0383	11.7	2.05	0.24
39	4.25	0.0512	0.0107	0.0137	4.05	2.36	0.245
40	3.86	0.0537	0.0103	0.0148	4.95	1.42	0.211
41	3.78	0.0711	0.0104	0.0194	6.60	1.56	0.222
42	3.59	0.0576	0.0107	0.0153	5.35	1.29	0.206
43	8.20	0.1040	0.0099	0.0300	4.53	6.22	0.324
44	6.05	0.0781	0.0092	0.0241	4.95	5.12	0.315
45	6.78	0.0823	0.0080	0.0292	5.36	5.70	0.323
46	6.23	0.0597	0.0081	0.0212	4.12	5.95	0.338
47	5.56	0.0815	0.0078	0.0300	6.60	5.10	0.320
48	4.51	0.0407	0.0073	0.0160	4.54	3.16	0.276
49	4.21	0.0496	0.0045	0.0318	9.45	2.07	0.237

INTAKE TOWER—8 meters

Sample No.	V	E	Δq	D
1	4.29	0.0740	0.0074	0.0287
2	4.73	-0.0157	0.0089	-0.0051
3	4.72	0.1141	0.0088	0.0370
4	7.41	0.1580	0.0083	0.0553
5	7.28	0.1040	0.0070	0.0432
6	8.08	0.1328	0.0043	0.0890
7	6.70	0.0844	0.0064	0.0375
8	6.85	0.0664	0.0072	0.0262
9	5.85	0.0604	0.0053	0.0328
10	4.99	0.0550	0.0042	0.0373
11	7.50	0.0653	0.0047	0.0394
12	5.95	0.0417	0.0042	0.0279
13	7.49	0.0366	0.0034	0.0305
14	4.69	0.0287	0.0033	0.0250
15	5.25	0.0059	0.0030	0.0057
16	6.75	0.0383	0.0037	0.0298
17	4.65	0.0182	0.0024	0.0218
18	5.42	0.0179	0.0029	0.0178
19	6.68	0.0426	0.0030	0.0405
20	6.56	0.0410	0.0034	0.0344
21	5.32	0.0222	0.0030	0.0211
22	5.08	0.0382	0.0035	0.0311
23	6.12	0.0295	0.0025	0.0337
24	5.74	0.0246	0.0024	0.0295
25	4.83	0.0076	0.0031	0.0071
26	8.67	0.0206	0.0044	0.0135
27	2.47	0.0052	0.0039	0.0036
28	5.90	0.0399	0.0042	0.0273
29	5.16	0.0056		
30	3.81	0.0182	0.0044	0.0118
31	3.99	0.0292	0.0038	0.0219
32	4.11	0.0192	0.0031	0.0177
33	2.35	0.0066	0.0035	0.00535
34	4.25	0.0253	0.0026	0.0280
35	3.47	-0.0874	0.0034	-0.0733
36	4.91	0.0005	0.0039	0.0004
37	6.73	0.0205	0.0037	0.0024
38	4.31	0.0311	0.0039	0.0227
39	4.74	0.0458	0.0100	0.0132
40	4.31	0.0794	0.0087	0.0259
41	4.32	0.0707	0.0089	0.0225
42	4.02	0.0758	0.0079	0.0272
43	9.17	0.0310	0.0072	0.0124
44	6.81	0.0891	0.0094	0.0269
45	7.72	0.0744	0.0083	0.0744
46	7.04	0.0798	0.0079	0.0798
47	6.28	0.1076	0.0076	0.1076
48	5.0	0.0671	0.0069	0.0671
49	4.50	0.0353	0.0070	0.0353

ABSTRACT

The evaporation rate from a water surface is shown to be greatly affected by surface waves in laboratory and field tests. Laboratory experiments were made using a large wave tank-wind tunnel apparatus in which both wind speed and wave parameters could be controlled. The apparatus simulates wave conditions on large bodies of water with fetches up to a mile or more. Tank experiments indicate certain combinations of wind speed and well-developed waves produce an evaporation rate lower than the rate under similar wind conditions with no waves present. Investigations of air flow over the waves show that, as wave heights increase, regions of dead air are observed leeward of wave crests, and vortices appear in the wave troughs. The partially trapped air apparently forms an effective barrier to the vertical transport of water vapor. Wave and evaporation measure-

ments at Lake Hefner, Oklahoma, were very similar to the laboratory measurements. The Lake Hefner data were adjusted to previously measured evaporation coefficient relationships derived from water and energy budget measurements. The resulting evaporation coefficient curve is believed to give an improved measure of the instantaneous evaporation rate from Lake Hefner under given conditions for wind speed, water surface temperature, and ambient humidity.

DESCRIPTORS—*evaporation/evaporation control / waves / *waves (water) / wind tunnels / wind velocity / wind speed / water vapor / water temperature / *reservoir evaporation / lakes / laboratory tests / field investigations.

IDENTIFIERS—*evaporation coefficient/evaporation tanks / Lake Hefner, Okla. / wave tanks / *wave height / mass transport.